

1 Experimental study of the nonlinear flow characteristics of fluid
2 in 3D rough-walled fractures during shear process

3 Changsheng Wang^{1,2}, Yujing Jiang^{1,2,*}, Richeng Liu,^{2,3} Chen Wang²,
4 Zhenyu Zhang⁴, Satoshi Sugimoto²

5 ¹ State Key Laboratory of Mining Disaster Prevention and Control Co-founded by
6 Shandong Province and the Ministry of Science and Technology, Shandong University of
7 Science and Technology, China

8 ² School of Engineering, Nagasaki University, Nagasaki 852-8521, Japan

9 ³ State Key Laboratory for Geomechanics and Deep Underground Engineering, China
10 University of Mining and Technology, Xuzhou 221116, P.R. China

11 ⁴State Key laboratory of Coal Mine Disaster Dynamics and Control, Chongqing
12 University, Chongqing 400044, P.R. China

13 Corresponding author: Yujing Jiang

14 E-mail address: jiang@nagasaki-u.ac.jp (Y. Jiang).

15 **Keywords:** Rough-walled fractures; shear-flow test; nonlinear flow; critical hydraulic
16 gradient; normalized transmissivity; visualization.

17
18 **List of symbols**

19	A	Cross-sectional area of a fracture
20	a	Linear coefficient in the Forchheimer's Law
21	b	Nonlinear coefficient in the Forchheimer's Law
22	b_h	Hydraulic aperture
23	d	Shear displacement
24	d_{contra}	Minimum normal displacement
25	d_{dilation}	Dilation of the fracture
26	d_{normal}	Ultimate normal displacement

27	d_{peak}	Peak shear displacement
28	E	Nonlinear effect factor
29	J	Hydraulic gradient
30	J_c	Critical hydraulic gradient
31	JRC	Joint roughness coefficient
32	k	Fracture permeability
33	K_s	Shear stiffness
34	L	Fracture length
35	P	Hydraulic pressure
36	Q	Volumetric flow rate
37	Re	Reynolds number
38	Re_c	Critical Reynolds number
39	T	Transmissivity
40	T/T_0	Normalized transmissivity
41	w	Width of a fracture
42	z_i	Coordinates of the fracture surface profile
43	Z_2	Dimensionless roughness parameter
44	ρ	Fluid density
45	μ	Dynamic viscosity
46	σ_n	Normal stress
47	τ_{peak}	Peak shear strength
48		

49 **Abstract:** To understand the influence of shear on the hydraulic properties of rock
50 fractures, shear-flow tests were carried out on rock fractures with different surface
51 roughnesses. Each rough-walled fracture was replicated in four specimens, which were
52 sheared at different displacements under normal stresses that varied from 0.5 MPa to 2.0
53 MPa. At each shear displacement, a series of hydraulic tests with different hydraulic
54 gradients were performed, and the nonlinear flow regimes of the fluid within the
55 fractures were investigated. The results show that Forchheimer's law can well describe
56 the nonlinear relationship between the flow rate and the hydraulic gradient in
57 rough-walled fractures. Both the linear coefficient and nonlinear coefficient decrease

58 during shearing but increase as the normal stress increases. The critical hydraulic
59 gradient increases with an increase in the shear displacement and normal stress. With an
60 increase in the joint roughness coefficient, the critical hydraulic gradient decreases. The
61 normalized transmissivity exhibits a strong correlation with the Reynolds number. As
62 the shear displacement increases, the fitted curves of the normalized transmissivity
63 versus the Reynolds number shift upward but the curves shift downward with an
64 increase in normal stress. Additionally, the Forchheimer coefficient decreases with an
65 increase in the shear displacement but increases with an increase in the applied normal
66 stress. Visualization tests show that the number of flow paths is large when the shear
67 displacement is small due to various distributions of the contact areas and that the flow
68 of dyed water over the entire fracture decreases. As the shear displacement increases,
69 the flow resistance decreases due to the shear dilation-induced increase in the aperture,
70 and the advantage channel flow is distinct in the fracture. The contact ratio rapidly
71 decreases as the shear displacement increases from 1 to 3 mm and then slightly varies
72 with a continuously increasing maximum shear displacement of 9 mm.

73 **1. Introduction**

74 Knowledge of fluid flow in fractured rock masses is critical to natural gas
75 production, geological sequestration of carbon dioxide, radioactive waste disposal and
76 geothermal energy extraction (Wu et al. 2011; Babadagli et al. 2015; Leung and
77 Zimmerman, 2012; Zhou et al. 2015; Wang et al. 2017; Chen et al. 2017). Rock masses
78 are a typical dual-porosity system that includes a rock matrix surrounded by fracture
79 networks. Fracture networks are considered the main channels for fluid flow due to their
80 substantially higher permeability compared with a rock matrix (Liu et al. 2011; Wang et

81 al. 2016; Wei and Xia 2017). Thus, understanding the fluid flow properties of rock
82 fractures is fundamental to the performance and safety assessment of these types of
83 engineering projects.

84 Individual fractures in natural rock masses form fracture networks; therefore,
85 understanding fluid flow in a single fracture is a foundation for modeling flow via
86 complex fracture networks (Rong et al. 2016). Generally, fluid flow in a single fracture
87 obeys the well-known Navier–Stokes (NS) equations (Zimmerman and Bodvarsson,
88 1996; Koyama et al. 2008). However, due to complex nonlinear partial differential
89 equations and irregular fracture morphology, solving the NS equations is difficult
90 (Brush and Thomson, 2003; Javadi et al.2010; Huang et al. 2017). To circumvent this
91 problem, for steady laminar flow through single fractures, the inertial terms of the NS
92 equations can be considered negligible, and the NS equations can be simplified to the
93 cubic law (Zhang and Nemcik 2013; Tzelepis et al. 2015; Li et al. 2016), which can be
94 written as:

$$95 \quad Q = -\frac{\Delta P}{L\mu} \frac{wb_h^3}{12} \quad (1)$$

96 where Q is the volumetric flow rate, ΔP is the pressure drop in the flow direction, w is
97 the fracture width, b_h is the hydraulic aperture, L is the fracture length over which the
98 pressure drop occurs, and μ is the viscosity of the fluid. The fracture transmissivity T is
99 equal to the term $wb_h^3/12$ in Eq. (1), which can be written as (Brown et al. 1995;
100 Olsson and Barton 2001):

$$101 \quad T = kA = \frac{wb_h^3}{12} \quad (2)$$

102 where $k = b_h^2/12$ and $A = wb_h$, where k is the permeability of the fracture and A is the
103 cross-sectional area. The cubic law shows a linear relationship between the flow rate

104 and the pressure drop and assumes that the fracture surface can be approximated by the
105 flat parallel plate model. However, the cubic law is derived from disregarding the inertia
106 effects, which is valid for laminar flow with a sufficiently low flow rate. In nature, the
107 rock fracture in realistic situations exhibits a complexity of the geometrical
108 characteristics on their surfaces, which causes deviations from the cubic law. In this
109 case, use of the cubic law to calculate fluid flow overestimates the transmissivity of
110 rock fractures (Zimmerman and Bodvarsson 1996; Oron and Berkowitz 1998; Wang et
111 al. 2015; Zhang et al. 2017; Yu et al. 2017). Therefore, many studies adopted
112 Forchheimer equations to describe the nonlinear flow through rough-walled rock
113 fractures (Forchheimer 1901):

$$114 \quad J = aQ + bQ^2 \quad (3)$$

$$115 \quad a = -\mu/\rho gkA, \quad b = -\beta\mu/\rho gA^2 \quad (4)$$

116 where $J = \Delta P/\rho gL$, J is the hydraulic gradient, ρ is the fluid density, g is the
117 gravitational acceleration and β is the Forchheimer coefficient. a and b are coefficients
118 that represent the pressure drops caused by linear effects and nonlinear effects,
119 respectively. Eq. (3) can well describe the nonlinear flow behaviors in fractures, which
120 has been verified to be valid via experiments, numerical simulations, and theoretical
121 analyses (Moutsopoulos 2009; Qian et al. 2011; Cherubini et al. 2012; Adler et al. 2013).
122 However, the mechanisms that trigger nonlinearity in rough-walled rock fractures are
123 not completely understood because this equation cannot consider the effects of the
124 fracture surface roughness and contact area on the flow behavior (Rong et al. 2017).
125 Chen et al. (2000) reported that the fracture surface roughness, contact area and
126 tortuosity may cause nonnegligible inertial forces that cause deviation from the cubic
127 law. Similar conclusions were obtained by Brush and Thomson (2003) and Li et al.

128 (2008), who concluded that complicated flow patterns cause inertial losses even with
129 laminar flow.

130 Previous studies have investigated many factors that can influence the nonlinear
131 flow characteristics in fractures, such as normal stress (Rabijth and Darlington 2007;
132 Zhang and Nemcik 2013; Zhou et al. 2015; Chen et al. 2015), fracture surface
133 roughness (Javadi et al. 2010; Wang et al.2016; Liu et al. 2016), fracture aperture
134 (Koyama et al .2006; Liu et al. 2016a) and fracture intersection (Kosakowski and
135 Berkowitz 1999; Liu et al. 2016a; Liu et al. 2016c, Yin et al. 2018).The critical
136 Reynolds number (Re_c) and critical hydraulic gradient (J_c) were commonly defined in
137 these studies to quantify the onset of nonlinear fluid flow. Rabijth and Darlington (2007)
138 observed that Re_c decreases from 31.29 to 28.77 with an incease in the confining stress
139 from 1.0 MPa to 5.0 MPa. Wang et al. (2016) discovered that Re_c decreases following
140 an exponential function as the fracture surface roughness increases. Liu et al. (2016b)
141 investigated the effect of the fracture aperture on the nonlinear flow regimes. The results
142 show that J_c decreases by approximately four orders of magnitude with an increase in
143 the fracture aperture from 1.0 mm to 10.0 mm. Kosakowski and Berkowitz (1999)
144 established numerical models with different single fracture intersection geometries to
145 study the validity of Darcy's law. The results indicate that the nonlinear inertial effects
146 become important at Reynolds numbers (Re) from 1 to 100. Liu et al. (2016a) estimated
147 the influence of the number of fracture intersections on the nonlinear flow in fracture
148 networks. The results show that J_c initially significantly decreases and then slowly
149 decreases with a power law function with an increase in the number of fracture
150 intersections from 1 to 12.

151 However, most previous studies are focused on flow characterizations of well-mated

152 fractures. Note that the fractures are commonly subjected to shear, which is caused by
153 various natural and human activities, such as earthquakes, underground excavations and
154 geothermal energy reservoir productions. Shear displacement can cause a change in the
155 contact area and the aperture distribution, which is an important issue that affects flow
156 and transport behavior in rough-walled fractures. (Yeo et al. 1998; Esaki et al. 1999; Yin
157 et al. 2017). Therefore, investigating the effects of the shear process on the nonlinear
158 flow characteristics of fractures is important. Javadi et al. (2014) conducted coupled
159 shear-flow tests on three granite specimens to investigate the variation in Re_c during the
160 shear process under different normal stresses. The results showed that Re_c ranges from
161 0.001 to 25 as the shear displacement increases from 0 to 20 mm. However, the effect of
162 the joint roughness coefficient on nonlinear flow regimes in fractures during shear was
163 negligible. Similar shear-flow tests were also conducted by Rong et al. (2016), who
164 performed shear-flow tests on six granite fractures with the joint roughness coefficient
165 (JRC) = 6.67 ~ 8.18 under different normal stresses to study the nonlinear fluid flow
166 characteristics during shear. They suggested that Re_c ranges from 1.5 to 13.0 as the
167 shear displacement increases from 0 to 10.9 mm. However, the range of JRC (8.18 –
168 6.67 = 1.51) was very small, and the effect of the JRC on the nonlinear fluid flow
169 behavior was not clearly presented. Recently, Yin et al. (2017) experimentally analyzed
170 the influences of the shear displacement and normal stress on the nonlinear flow
171 behavior of fluid through 3D rough-walled rock fractures with JRC = 15.17 and
172 discovered that J_c increased with an increase in the shear displacement and decreased
173 with an increase in the normal stress. As previously described, most previous studies
174 have focused on the effect of the normal stress on the nonlinear fluid flow
175 characteristics during shear. However, few studies have focused on the influence of the

176 fracture surface roughness on the characteristics of nonlinear fluid flow through
177 fractures during shear.

178 This study systematically investigates the effects of the fracture surface roughness,
179 normal stress and shear displacement on the nonlinear flow characteristics of a single
180 fracture. Four types of granite fractures with different surface characteristics were
181 employed in the tests. At each shear step, high precision hydraulic tests were conducted
182 with different hydraulic gradients. The nonlinear flow regimes are analyzed in terms of
183 the fracture surface roughness and normal stress of fractures during shear. Additionally,
184 to estimate the influence of the shear process on the nonlinear flow characteristics in
185 rough-walled fractures, a visualization test was conducted on fracture G3-1 to determine
186 how the channeling flow and contact area affect the nonlinear flow regimes.

187 **2. Experimental method**

188 **2.1 Specimen preparation and roughness measurement**

189 In this study, four granite specimens were collected from the Omarugawa power
190 station in Miyazaki prefecture, Japan. The four specimens were cut and polished to
191 cuboid specimens with the following dimensions: length of 200 mm, width of 100 mm
192 and height of 100 mm. An artificial tensile fracture was created using the Brazilian test.
193 During the test, a normal stress of 10 kN was applied to the specimen, and lateral stress
194 was then applied using V-shaped wedges to a constant horizontal stress of 120 kN. The
195 split was extended until a tensile fracture formed. Four fractures (labeled G1, G2, G3
196 and G4) with different surface geometries are shown in Fig. 1. One half of each
197 specimen was chosen as the model to cast another half using resin material. The upper

198 and lower halves of the specimens were then manufactured based on the resin replica.
199 For visual shear-flow tests, a transparent acrylic specimen was prepared as the upper
200 half, and the plaster specimen was fixed as the lower half. Each rough-walled fracture
201 was replicated in four specimens. The artificial fracture specimens were manufactured
202 using mixtures of plaster, water and retardant with a weight ratio of 1: 0.2: 0.005. The
203 physico-mechanical properties of these rock-like specimens are shown in Table 1.

204 The four fracture surfaces of the specimens were scanned at resolutions of $\pm 20 \mu\text{m}$
205 and $\pm 10 \mu\text{m}$ in the x - y plane and on the z axis, respectively. Based on the scanned data,
206 digitized fracture surfaces of the four specimens were created (Fig. 1). To quantify the
207 surface roughness of the fracture, ten profiles parallel to the flow direction were
208 extracted from each fracture by dividing the fracture into nine equal areas. The JRC was
209 calculated according to the equations proposed by Tse and Cruden (1979) and written as

$$210 \quad Z_2 = \left[\frac{1}{(n-1)(\Delta x)^2} \sum_{i=1}^{n-1} (z_{i+1} - z_i)^2 \right]^{1/2} \quad (5)$$

$$211 \quad \text{JRC} = 32.2 + 32.47 \log Z_2 \quad (6)$$

212 where Z_2 is the root mean square slope of the profiles based on the extracted data, z_i
213 represents the coordinates of the fracture surface profile, n is the number of data points,
214 and Δx is the interval between the data points. The mean JRC values of fractures G1,
215 G2, G3 and G4 were 3.21, 5.62, 7.36 and 12.16, respectively. The variation in frequency
216 versus asperity height of four fracture surfaces are depicted in Fig. 2; the results show
217 that the rougher fracture trends to generate a wider range of the asperities distribution.

218 2.2 Experimental apparatus

219 In this study, shear-flow tests were carried out using a laboratory visualization

220 system of the shear-flow test apparatus, as shown in Fig. 3. The system mainly consists
221 of a hydraulic-servo actuator unit, a hydraulic testing unit and a visualization unit. The
222 hydraulic-servo actuator unit consists of two load jacks that apply the normal and shear
223 loads via a servo-controlled hydraulic pump. The maximum applied load is 200 kN in
224 the shear and normal direction with a precision of 99%. The hydraulic testing unit
225 contains a water supply system, a fracture sealing system and a measurement system.
226 Constant water pressure was applied to one side of the fracture using an air pump. A
227 pressure transducer with a precision of 0.01 kPa was attached to the water inlet to
228 measure the inlet water pressure. The water outlet was connected to a tube with a
229 diameter of 10 mm and a smooth inner wall. The outlet pressure was assumed to be zero
230 due to the relatively small flow rate. The lateral sides of the fracture were sealed using
231 soft and elastic gel sheets, which were flexible and strong with a good sealing effect.
232 The water that flows out of the fracture was measured using an electrical balance with a
233 precision of 0.01 g. The most important characteristic of this apparatus is that an
234 observation window is created above the upper part of the shear box and a normal load
235 is applied to the specimen from the bottom of the shear-box. Hence, a coupled charged
236 device (CCD) camera placed on top of the specimen can directly view the dyed water
237 that flows through the transparent acrylic specimen. Many coupled shear-flow tests have
238 been carried out using this apparatus, which shows a good sealing effect (i.e., Li et al.
239 2008; Koyama et al. 2008; Xiong et al. 2011).

240 2.3 Experimental procedures

241 A specimen was set in the shear box, and the shear box was fixed on the steel plate
242 connected to the normal load jack. The lower shear box can only move in the vertical

243 direction by the roller guide, and the upper shear box can move in the vertical and
244 horizontal direction without rotation during shearing (Xiong et al. 2011). Two small
245 tanks were fixed at the inlet and outlet of the specimen, which were sealed using rubber
246 sheets on the two ends of the specimen. The lateral sides were sealed using soft and
247 elastic gel sheets. Coupled shear-flow tests were conducted at five shear displacements
248 d (1.0, 3.0, 5.0, 7.0 and 9.0 mm) under constant normal stresses σ_n of 0.5, 1.0, 1.5 and
249 2.0 MPa. The test cases and their corresponding normal stresses are shown in Table 2.
250 The shear velocity was 0.5 mm/min. At each shear step, 7 ~ 10 constant head water flow
251 tests were carried out. The constant water pressure was supplied using an air pump. The
252 flow rate was measured by collecting the discharge from the outlet with a high precision
253 electrical balance. Shear-flow tests were performed at a room temperature of
254 approximately 20 °C. The density and dynamic viscosity of water are $\rho = 0.998 \times 10^3$
255 kg/m³ and $\mu = 1.006 \times 10^{-3}$ Pa·s, respectively.

256 **3. Experimental results and discussions**

257 **3.1 Mechanical behavior**

258 Sixteen replicated fracture specimens with four different fracture surface
259 roughnesses were used to conduct coupled shear-flow tests under various normal
260 stresses. The mechanical shear behaviors and characteristics of the fractures are shown
261 in Fig. 4, Fig. 5 and Table 2. The data clearly show that the shear stress abruptly
262 increases to the peak at the very beginning of shear. Then, the shear stress gradually
263 decreases to the residual stage. The peak shear displacement, d_{peak} , ranges from 0.844
264 mm to 1.428 mm. The peak shear strength, τ_{peak} , ranges from 0.515 MPa to 1.789 MPa

265 for G1 (Fig. 4 (a)), from 0.652 MPa to 1.834 MPa for G2 (Fig. 4 (b)), from 0.770 MPa
266 to 2.070 MPa for G3 (Fig. 4 (c)), and from 0.813 MPa to 2.141 MPa for G4 (Fig. 4 (d)).
267 The peak shear strength increases with an increase in the normal stress and JRC. The
268 influences of normal stress and JRC on shear stiffness (K_s) is plotted in Fig. 6. Shear
269 stiffness is defined as the slope of the pre-peak stage of the shear stress-shear
270 displacement curves. As shown in Fig. 6, K_s shows an increasing trend with an
271 increasing normal stress and JRC.

272 The normal displacement is an important parameter in shear-flow tests for
273 quantifying the permeability of fractures due to increases or decreases in the fracture
274 aperture. As shown in Fig. 7, the normal displacement decreases at the onset of shear
275 and then rapidly increases; however, the increasing rate decreases. The decrease in
276 normal displacement is due to the deformation of asperities and surface interlocking.
277 The minimum normal displacement, d_{contra} , ranges from -0.03 mm to -0.172 mm for all
278 test cases and depends on the applied normal stress. The larger is σ_n , the larger is d_{contra} .
279 With continuously increasing shear displacement, the normal displacement increases
280 due to shear dilation. The dilation causes the contact area to significantly decrease due
281 to the overriding and sliding of the contact asperities. At this stage, the aperture between
282 the opposite surfaces rapidly increases, which causes the permeability to substantially
283 increase (Rong et al. 2016). The dilation increasing rate decreases by abrading some
284 rougher asperities until the residual stage is reached. The ultimate normal displacement
285 d_{normal} (corresponds to a shear displacement of 9.0 mm), ranges from 0.560 mm to 2.467
286 mm as listed in Table 2. The results show that d_{normal} increases with increasing normal
287 stress, and the rougher fracture shows a larger dilation in the same test conditions.

288 The shear stress and normal deformation during the shear process exhibit a

289 three-stage behavior (Xiong et al. 2011). First, the shear stress rapidly increases to the
290 peak value, while the normal displacement reaches the maximum negative value.
291 Second, the shear stress gradually decreases, while the normal displacement rapidly
292 increases. Last, the shear stress and normal displacement reach the residual values in the
293 third stage, during which the rate of decrease of the shear stress decreases, and the
294 normal displacement continues to increase at a lower gradient. Our test results have
295 similar trends to those in previous studies (Li et al. 2008; Koyama et al. 2008; Xiong et
296 al. 2011; Javadi et al. 2014; Rong et al. 2016).

297 3.2 Fluid flow behavior

298 3.2.1 Effect of normal stress

299 In this section, eighty individual hydraulic tests based on the sixteen specimens were
300 conducted with $d = 1, 3, 5, 7$ and 9 mm. To avoid the influence of gouge materials on
301 the fluid flow behavior during the shear process, the shear-flow tests were conducted at
302 a small range of normal stresses ($0.5 \sim 2.0$ MPa). During the tests, only a small amount
303 of gouge material was observed. Therefore, the effect of the gouge material on the fluid
304 flow is negligible.

305 Fig. 7 shows the relationships between the hydraulic gradient J and the flow rate Q ,
306 which corresponds to different shear displacement for fractures G1 at different constant
307 values of σ_n (other test data sets are attached in “Appendix A, Fig. A1”). The
308 relationship between J and Q exhibits distinct nonlinear characteristics. The
309 Forchheimer equation fits the data well with a residual squared (R^2) greater than 0.98 in
310 all cases. With an increase in d , the slopes of the $J - Q$ curves become flatter as a result
311 of fracture dilation (d_{dilation}) during shear. At a given shear displacement, as σ_n increases,

312 Q decreases with nearly the same hydraulic gradient. For a higher σ_n , this phenomenon
313 indicates that a larger hydraulic gradient is required to achieve the same flow rate. Note
314 that the water flowed from a relatively large tank on the sealed side boundary into the
315 relatively small void space of the fractures during the hydraulic tests. With a high
316 hydraulic gradient, the fluid may flow back and influence the nonlinear flow
317 characteristics. However, this is difficult to avoid using the current techniques. In future
318 studies, we will attempt to decrease this system error.

319 Forchheimer's law is the most extensively employed mathematical description of
320 the nonlinear flow in fractures (Zimmerman et al. 2004; Zhang and Nemcik 2013;
321 Javadi et al. 2014; Wang et al. 2016; Rong et al. 2016; Yin et al. 2017). Based on Eq. (3),
322 the linear coefficient a and nonlinear coefficients b were calculated, as plotted in Fig. 8
323 and Fig. 9, respectively. Both a and b decrease with an increase in d , and the decrease
324 extent at d from 1 to 3 mm is more significant than the reduction in the range from 3 to
325 9 mm. The decrease in a with d from 1 to 9 mm was approximately 0.5 ~ 1.0 orders of
326 magnitude, and the decrease in b was approximately 2.0 ~ 2.5 orders of magnitude.
327 According to Eq. (4), the coefficient a is related to the permeability of the fracture, and
328 a lower value of a means a higher value of permeability. Therefore, the decrease in the
329 linear coefficient a is caused by dilation of the fracture aperture that increases the
330 fracture permeability. As d increases from 1 mm to 3 mm, the coefficient a significantly
331 decreases, and then the decreasing rate gradually decreases. This conclusion is
332 consistent with the experimental findings in this study (Javadi et al. 2014; Rong et al.
333 2016 and Yin et al. 2017). The coefficient a increases with an increase in the normal
334 stress. This result is primarily attributed to the closure of the fracture in higher normal
335 stress conditions. Similarly, according to Eq. (4), the decrease in b is attributed to the

336 shear dilation-induced increase in the fracture aperture and a decrease in the flow
 337 tortuosity caused by the decrease in the contact area. With the increment of confining
 338 stress, the increase in b is ascribed to the closure of the fracture and an increase in the
 339 flow path tortuosity as the contact area increases. These phenomena can be verified by
 340 the relationship between a and b with d_{dilation} . As shown in Fig. 8 and Fig. 9, both the
 341 coefficient a and coefficient b show a decreasing trend with an increase in d_{dilation} .
 342 Additionally, the coefficient a shows a decreasing trend with an increase in JRC due to
 343 the rougher fractures that cause a larger dilation with the same d . However, the
 344 correction between coefficient b and JRC was not established in this study because the
 345 variation in coefficient b is controlled by the fracture roughness and fracture aperture.
 346 The rougher fracture surfaces cause increases in coefficient b , while the rougher
 347 fractures creates a larger aperture during the shear process, which causes a decrease in
 348 coefficient b . These two competitive effects produce a trend that was hard to establish
 349 for coefficient b and the fracture roughness.

350 3.2.2 Critical hydraulic gradient analysis

351 The nonlinear flow effects become more significant in fluid flow in fractures with
 352 an increasing flow velocity. The factor E has been used to determine the fluid flow
 353 regime (Zeng and Grigg 2006; Zhang and Nemeik 2013; Xia et al. 2016; Liu et al.
 354 2016), which can be written as

$$355 \quad E = \frac{bQ^2}{aQ + bQ^2} \quad (7)$$

356 where aQ and bQ^2 are energy losses due to the linear and nonlinear dissipation
 357 mechanisms in the fracture, respectively. E denotes the percentage of the nonlinear term
 358 that contributes to the ratio of the nonlinear term-induced decrease in the hydraulic
 359 gradient to the total decrease in hydraulic gradient. Based on Eq. (3), Q can be solved,

360 written as

$$361 \quad Q = \frac{aE}{b(1-E)} \quad (8)$$

362 The critical hydraulic gradient can be obtained by substituting Eq. (8) into Eq. (3),
363 which is written as follows:

$$364 \quad J_c = \frac{a^2E}{b(1-E)^2} \quad (9)$$

365 Generally, the critical condition for the transition of linear flow regimes to nonlinear
366 flow regimes has been defined as $E = 0.1$ (Zimmerman et al. 2004; Zhang and Nemcik
367 2013; Rong et al. 2016). The corresponding J is defined as the critical hydraulic
368 gradient J_c (Liu et al. 2016a; Liu et al. 2016b; Wang et al. 2016). In previous studies of
369 the flow characteristics of fractures, Re has been extensively employed to quantify the
370 nonlinear flow behavior, which can be written as follows (Ranjith and Darlington 2007;
371 Zhang and Nemcik 2013):

$$372 \quad Re = \frac{\rho Q}{\mu w} \quad (10)$$

373 In engineering practices, note that rock masses contain hundreds to thousands of
374 fractures, and the Re_c value of each fracture cannot be ascertained. However, the J_c
375 value can be easily obtained. In this study, therefore, J_c was used to evaluate the fracture
376 flow characteristics during shear (Liu et al. 2016a; Wang et al. 2016).

377 Fig. 10 shows the variation in J_c during shear with different σ_n for four different JRC
378 values. The results show that J_c increases and exhibits two stages as d increases. As d
379 increases from 1 mm to 3 mm, J_c changes significantly, and then the rate of increase
380 gradually decreases. The increase in J_c is mainly attributed to shear-induced dilation. As
381 shown in Fig. 10(e) ~ (h), J_c increases with an increase in d_{dilation} . Taking JRC = 5.62 as
382 an example, when $\sigma_n = 0.5$ MPa and the dilation increases from 0.026 mm to 0.504

383 mm, J_c increases from 0.98 to 3.42 and then increases from 3.42 to 4.36 with an
384 increase in dilation from 0.504 mm to 1.309 mm. This result occurs because the shear
385 dilation causes a distinct change in the fracture geometry and a decrease in the contact
386 ratio, relative roughness and flow tortuosity. These features decrease the inertial losses
387 and generate a large J_c . These results are generally consistent with the work by Yin et al.
388 (2017). In all cases, as σ_n increases, J_c increases. When $JRC = 12.16$ and d increases
389 from 1 to 9 mm, J_c varies from 0.29 to 1.31 for $\sigma_n = 0.5$ MPa and varies from 0.85 to
390 2.03 for $\sigma_n = 2.0$ MPa. This is mainly due to the closure of the fracture and increasing
391 contact areas caused by the higher σ_n . According to Eq. (1), Q is proportional to b_h^3 ;
392 thus, a slight decrease in the fracture aperture causes a large decrease in the flow rate.
393 Therefore, J_c increases with an increase in σ_n . The experimental results also show that
394 the JRC impacts the range of J_c subjected to a certain σ_n . The larger is JRC, the lower is
395 the range of J_c . When $\sigma_n = 1.0$ MPa and d increases from 1 mm to 9 mm, J_c ranges
396 from 1.21 to 8.19 for $JRC = 3.21$, ranges from 1.16 to 4.7 for $JRC = 5.62$, ranges from
397 0.74 to 4.38 for $JRC = 7.36$, and from 0.45 to 1.65 for $JRC = 12.16$. According to
398 Schrauf and Evans (1986), the occurrence of nonlinear flow is mainly attributed to the
399 inertial losses due to changes in the flow velocity of the direction along the flow paths.
400 Sharp changes in the fracture aperture along the flow path will promote the variations in
401 the plane velocity due to the acceleration and deceleration of flow to maintain the
402 conservation of mass. This acceleration and deceleration of flow cause deviation from a
403 linear relationship between pressure drop and flow rate (Cornwell and Murphy 1985;
404 Javadi et al. 2014; Wang et al. 2016). Therefore, the nonlinear flow is strongly related to
405 the aperture fields and fracture surface roughness. When shear displacement occurs, the
406 mismatch of the fracture surface causes a relative distribution of the asperities and

407 variable aperture structure (aperture field). The flow path is controlled by the
408 distribution of the aperture field together with the relative distribution of the asperities.
409 As shown in Fig. 2, the rougher fracture trends to generate a wider range of the
410 asperities distribution. This phenomenon indicates that the aperture distribution
411 becomes more anisotropic and heterogeneous for rougher fractures during the shear
412 process, which enhances the local complexity of the flow velocity distributions or
413 direction along the flow channel and enables the fluid flow to more easily become
414 nonlinear. In addition, the dilation of fractures for the same d is larger in the fractures
415 with a larger JRC value. The increase in the fracture aperture yields a lower hydraulic
416 gradient, which is required to obtain the same flow volume. Therefore, a lower value of
417 J_c is obtained for rougher fractures.

418 3.2.3 Normalized transmissivity

419 For fluid flow in fractures, transmissivity is an important parameter for estimating
420 the flow characteristics (Zimmerman et al. 2004; Wang et al. 2016). When the flow rate
421 is sufficiently low, the intrinsic transmissivity (T_0) is regarded a constant that is
422 independent of the flow rate. As the flow rate increases, the apparent transmissivity T
423 calculated using Eq. (2) can be applied to assess the nonlinear flow. The normalized
424 transmissivity (T/T_0) is defined as the ratio of the apparent transmissivity to the intrinsic
425 transmissivity, which can be described as a function of Re and the Forchheimer
426 coefficient β as follows (Zimmerman et al. 2004):

$$427 \quad \frac{T}{T_0} = \frac{1}{1+\beta Re} \quad (11)$$

428 where T_0 is determined using the best-fit values of the coefficient a .

429 The relationships between T/T_0 and Re for fractures with different surface
430 roughnesses are shown in Figs. 11 ~ 14. T/T_0 decreases with an increase in Re , which

431 further confirms the deviation of the flow from linearity. For a certain σ_n , as d increases,
432 the transmissivity relationship generally shifts upward. The shift degree of
433 transmissivity increases more significantly for d from 1 mm to 3 mm than from 3 mm to
434 9 mm. For a certain d , however, the relationship between Re and T/T_0 shifts downward
435 as σ_n increases.

436 Re_c was determined by establishing the fitting relations in the form of Eq. (9), as
437 shown in Table 3. Note that $T/T_0 = 0.9$ has the same physical meaning as $E = 0.1$,
438 which indicates that the nonlinear term contributes to 10% of the pressure drop. In this
439 case, the corresponding Re is defined as Re_c (Yu et al. 2017). The range of Re_c is 1.19 to
440 48.73 for all tested cases, which is the same range reported in the literature (Javadi et al.
441 2014; Rong et al. 2016).

442 The values of β for all cases are listed in Table 3. As d increases, β decreases.
443 Taking JRC = 12.16 and $\sigma_n = 1.0$ MPa as an example, with an increase in d from 1 to
444 9 mm, β decreases from 0.0823 to 0.00407. For a certain d , β generally increases
445 with an increase in σ_n . For JRC = 7.36 at $d = 5$ mm, β is 0.00456 for $\sigma_n = 0.5$ MPa,
446 0.00559 for $\sigma_n = 1.0$ MPa, 0.00642 for $\sigma_n = 1.5$ MPa, and 0.00859 for $\sigma_n = 2.0$ MPa.
447 However, the increasing trend was not distinct for $d = 1$ mm. This phenomenon occurs
448 because the shear contraction at $d = 1$ mm does not have a clear decreasing trend with
449 an increase in σ_n . It is also observed that β exhibits an increasing trend with an
450 increase in the JRC value. When $\sigma_n = 1.0$ MPa for JRC = 3.21, as d increases from 3
451 to 9 mm, β decreases from 0.00528 to 0.0025, while for JRC = 12.16, β decreases
452 from 0.01191 to 0.00407.

453 3.3 Visualization

454 Visual techniques have the advantage of directly observing the detailed flow
455 behaviors in rough-walled fractures. To provide a better understanding of the effect of
456 the shear process on the nonlinear flow regimes of rough-walled fractures, a
457 visualization shear-flow test was conducted on fracture G3 with a σ_n of 0.5 MPa,
458 which is denoted as fracture G3-1 in the following section. For visualization purposes,
459 the upper part of the rock fracture specimen was a transparent replica, and the lower part
460 of the rock fracture specimen was a plaster replica. The physic-mechanical properties of
461 the transparent acrylic material are shown in Table 1. Compared with the plaster
462 material, the transparent acrylic material has a larger value of uniaxial compressive
463 strength and a fairly similar value of elastic modulus. Shear displacement causes
464 degradation of fracture asperities, and different types of materials show different
465 mechanical and deformation behaviors during the shear process. This phenomenon
466 results in the different gouge productions and aperture fields during shear process and
467 influences the fluid flow behaviors in fractures. To avoid the influence of the material
468 property difference on fluid flow behavior, the visualization test was conducted at a
469 small value of normal stress (0.5 MPa). In this condition, the dilation of the transparent
470 specimen is considered to be fairly close to the plaster specimen because no distinct
471 gouge materials are generated during shearing. Since the visualization test mainly
472 focuses on the observation of fluid flow in the fracture during shearing, which is
473 primarily governed by the aperture changes induced by dilation, the acrylic-plaster pair
474 is considered an acceptable replica for studying the channelization characteristics of
475 fractures. Red-dyed water was used to enhance the visibility of the flow paths induced

476 by the shear process, whereas the effect of the shear-induced aperture field distribution
477 and tortuosity on nonlinear flow behavior were not quantitative analyzed due to the
478 limitation of the tracer technique. The visualization test was conducted with a constant
479 water head of 10 cm. This small inlet water head is used to ensure a small value of Re ;
480 thus, the channelized flow is generated by the shear-induced anisotropic void spaces
481 within the fractures rather than the inertial effects caused by a large water head
482 difference. A high-resolution CCD camera was placed on top of the shear box, which
483 can capture the fluid flow in the fracture in real time. The capture rate was one image
484 per second, which was maintained until the red-dyed water filled the void spaces in the
485 fracture. To maintain a constant starting point of the flow of water into the fracture, the
486 water inlet is set at the bottom of the tank (Fig. 3(b)). In the captured images, the area
487 invaded by the dyed water appears red, and the contact areas were easily distinguished
488 due to their yellow color. Therefore, the binary image method was used to calculate the
489 contact ratio (C) during different shear displacements. A detailed description of this
490 method is provided in Develi and Babadagli 2015.

491 Images of the fluid flow in fracture G3-1 at different shear displacements are shown
492 in Fig. 15. The channelized flow was observed at all shear steps. When the shear
493 displacement was small (i.e., $d = 1$ mm), the number of flow paths was larger due to the
494 widely distributed contact area, and the flow pattern was in the form of a network. The
495 dyed water slowly flowed over the entire fracture, which required 105 seconds to fill all
496 void spaces because the rough fracture surfaces may interlock with each other and form
497 a tightly closed fracture with a small aperture at very beginning of shear. As d increases,
498 the number of flow paths decreases due to the shear dilation-induced decrease in the
499 number of contact spots, and the preferential flow channels can be more

500 characteristically observed in the fracture plane. Additionally, the dyed water more
501 quickly reaches the stationary state at a larger shear displacement because the larger
502 aperture decreases the flow resistance. The same visualization test was conducted on
503 another fracture (G3) with a σ_n of 0.5 MPa. The visualization results show that the
504 preferential flow channels are similar, which indicates that the flow paths are repeatable
505 for tests at the same d . The variations in b_h , d_{dilation} and C at different shear
506 displacements are shown in Fig. 16(a). As d increases from 1 to 3 mm, b_h substantially
507 increases. When d ranges from 3 ~ 9 mm, b_h decreases and becomes steadier with an
508 increase in d mainly due to the dilation behavior of the fracture. As shown in Fig. 16(b),
509 the dilation significantly changes and then gradually decreases with an increase in d .
510 Additionally, the increase in d_{dilation} is greater than the b_h . This result shows agreement
511 with the observations of Olsson and Barton (2001). However, the C inversely changes
512 with a change in b_h during shear. As d increases from 1 to 3 mm, C rapidly decreases
513 from 32.9% to 9.27%. When d ranges from 3 ~ 9 mm, the change in C is very small,
514 and a nearly constant value is maintained. The roughness and contact area causes
515 tortuous flow paths, which can cause nonnegligible inertial losses. These features may
516 contribute to nonlinear fluid flow behavior. As shear begins, the flow path is controlled
517 by the relative distribution of the asperities and the distribution of the aperture field.
518 When d is very small (i.e., $d = 1$ mm), the opposite walls of a rough fracture are slightly
519 mismatched, and a void space with many contact spots forms. At this stage, the
520 roughness of the fracture surface has a dominant role in controlling the fluid flow, and
521 the existing large number of contact areas produce more tortuous flow paths. This
522 phenomenon causes an increase in the degree of the nonlinearity of the fluid flow and
523 decreases in Re_c and J_c . As the shear displacement increases (i.e., $d = 1 \sim 3$ mm), the

524 rapid increase in fracture aperture causes a rapid decrease in the number of contact areas.
525 Therefore, the influence of the flow tortuosity decreases. These physical processes
526 cause considerable increases in both Re_c and J_c . With a continuous increase in d to 9
527 mm, the variation in C is very small, and a nearly constant value is maintained.
528 Therefore, the effect of the contact-induced flow path tortuosity on nonlinear flow is not
529 the key factor. The influence of the fracture surface roughness on the nonlinear flow
530 decreases, and the shear dilation-induced increase in fracture aperture plays a dominant
531 role in controlling the nonlinear flow behavior, which causes gradual increases in Re_c
532 and J_c .

533 **4. Conclusions**

534 1) Forchheimer's law can well describe the nonlinear flow in rough-walled
535 fractures during shear at different normal stresses. Both the linear coefficient a
536 and the nonlinear coefficient b decrease as the shear dilation increases. The
537 degrees of the decreases in the two coefficients at the shear displacement of 1 to
538 3 mm are more significant than that for the shear displacement of 3 to 9 mm.
539 The decrease in a as d increased from 1 to 9 mm was approximately 0.5 ~ 1.0
540 orders of magnitude, and the decrease in b was approximately 2.0 ~ 2.5 orders
541 of magnitude. The coefficients a and b are very sensitive to the normal stress.
542 Both a and b decrease with an increase in the normal stress.

543 2) With an increase in shear dilation, the critical hydraulic gradient increases in
544 two different stages. As the shear displacement increases from 1 mm to 3 mm,
545 the critical hydraulic gradient significantly changes, and the rate of increase
546 then gradually decreases with an increase in the shear displacement. As normal

547 stress increases, the critical hydraulic gradient increases due to the fracture
548 closure. In addition, the critical hydraulic gradient is influenced by the
549 roughness of the fracture surfaces. With an increase in the fracture surface
550 roughness, the critical hydraulic gradient decreases.

551 3) The normalized transmissivity, which is a function of the hydraulic gradient, is
552 analyzed. As the shear displacement increases, the fitted curves shift upward
553 and then downward with an increase in normal load. The coefficient β decreases
554 with an increase in shear displacement but increases with an increase in normal
555 load. The value of β increases with an increase in fracture surface roughness.

556 4) The visualization tests show that channelized flow occurs in all cases. When the
557 shear displacement is small, the number of flow paths is large due to the
558 distribution of the contact areas, and the dyed water slowly flows over the entire
559 fracture. With an increase in shear displacement, the shear dilation-induced
560 increase in fracture aperture causes the contact ratio to rapidly decrease to a
561 small value and then maintain an approximately constant value. The flow paths
562 are focused in the void spaces with relatively large apertures, in which some
563 preferential flow channels form. Channels with larger apertures have lower flow
564 resistance, and the dyed water can reach the stationary state faster with an
565 increase in shear displacement.

566 In this study, we investigated the effect of shear on the nonlinearity of fluid flow in
567 single rough-walled rock fractures. However, more in-depth research of this issue is
568 required. In future studies, we will focus on the effects of the fracture surface roughness
569 and shear behavior on the nonlinear flow regime in complex fracture networks.

570 **Acknowledgments**

571 This study has been partially funded by National Natural Science Foundation of
572 China (Grant Nos. 51979272 and 51709260), Natural Science Foundation of Jiangsu
573 Province, China (Grant No. BK20170276), JSPS-NSFC Bilateral Joint Research Project
574 (Grant No. 51611140122) and China Scholarship Council (CSC NO. 201608370099).
575 These supports are gratefully acknowledged.

576 **Conflict of Interest**

577 The authors declare that they have no conflict of interest.

578 **References**

- 579 Adler PM, Malevich AE, Mityushev VV. Nonlinear correction to Darcy's law for
580 channels with wavy walls. *Acta Mechanica*, 2013, 224(8):1823-1848.
- 581 Brown S R. Fluid flow through rock joints: The effect of surface roughness. *Journal of*
582 *Geophysical Research Solid Earth*, 1987, 92(B2):1337-1347.
- 583 Brown S R, Stockman H W, Reeves S J. Applicability of the Reynolds Equation for
584 modeling fluid flow between rough surfaces. *Geophysical Research Letters*, 1995,
585 22(18):2537-2540.
- 586 Brush DJ, Thomson NR. Fluid flow in synthetic rough-walled fractures: Navier-Stokes,
587 Stokes, and local cubic law simulations. *Water Resources Research*, 2011,
588 39(4):1037-1041.
- 589 Babadagli T, Ren XJ, Develi K. Effects of fractal surface roughness and lithology on
590 single and multiphase flow in a single fracture: an experimental investigation.

591 International Journal of Multiphase Flow. 2015, 68:40–58.

592 Chen YD, Liang WG, Lian HJ, Yang JF, Nguyen VP. Experimental study on the effect of
593 fracture geometric characteristics on the permeability in deformable rough-walled
594 fractures. International Journal of Rock Mechanics & Mining Sciences, 2017,
595 98:121-140.

596 Chen YF, Hu SH, Hu R, Zhou CB. Estimating hydraulic conductivity of fractured rocks
597 from high - pressure packer tests with an Izbash's law - based empirical model.
598 Water Resources Research, 2015, 51(4):2096-2118.

599 Cherubini C, Giasi CI, Pastore N. Bench scale laboratory tests to analyze non-linear flow
600 in fractured media. Hydrology & Earth System Sciences Discussions, 2012,
601 9(4):2511-2522.

602 Chen Z, Narayan SP, Yang Z, Rahman SS. An experimental investigation of hydraulic
603 behaviour of fractures and joints in granitic rock. International Journal of Rock
604 Mechanics & Mining Sciences 2000; 37:1061–71.

605 Cornwell, D.K, Murphy, et al. Experiments with non-Darcy flow in joints with large
606 scale roughness. International Symposium on Fundamentals of Rock Joints, 1985,
607 Centek, Lulea.

608 Develi K, Babadagli T. Experimental and visual analysis of single-phase flow through
609 rough fracture replicas. International Journal of Rock Mechanics & Mining Sciences,
610 2015, 73:139-155.

611 Esaki T, Du S, Mitani Y, Ikusada, Jing L. Development of shear-flow test apparatus and
612 determination of coupled properties for a single rock joint. International Journal of
613 Rock Mechanics & Mining Sciences, 1999, 36(5):641-650.

614 Huang N, Liu R, Jiang Y. Numerical study of the geometrical and hydraulic

615 characteristics of 3D self-affine rough fractures during shear. *Journal of Natural Gas*
616 *Science & Engineering*, 2017, 45.

617 Javadi M, Sharifzadeh M, Shahriar K. A new geometrical model for non-linear fluid flow
618 through rough fractures. *Journal of Hydrology*, 2010, 389(1):18-30.

619 Javadi M, Sharifzadeh M, Shahriar K, Mitani Y. Critical Reynolds number for nonlinear
620 flow through rough-walled fractures: The role of shear processes. *Water Resources*
621 *Research*, 2014, 50(2):1789–1804.

622 Kosakowski G, Berkowitz B. Flow pattern variability in natural fracture intersections.
623 *Geophysical Research Letters*, 1999, 26(12):1765–1768.

624 Koyama T, Li B, Jiang YJ, Jing LR. Coupled shear-flow tests for rock fractures with
625 visualization of the fluid flow and their numerical simulations. *International Journal of*
626 *Geotechnical Engineering*, 2012, 2(3):215-227.

627 Koyama T, Neretnieks I, Jing L. A numerical study on differences in using Navier–Stokes
628 and Reynolds equations for modeling the fluid flow and particle transport in single
629 rock fractures with shear. *International Journal of Rock Mechanics & Mining Sciences*,
630 2008, 45(7):1082-1101.

631 Li B, Jiang YJ, Koyama T, Jing LR. Experimental study of the hydro-mechanical
632 behavior of rock joints using a parallel-plate model containing contact areas and
633 artificial fractures. *International Journal of Rock Mechanics & Mining Sciences*, 2008,
634 45(3):362-375.

635 Li B, Liu R, Jiang Y J. Influences of hydraulic gradient, surface roughness, intersecting
636 angle, and scale effect on nonlinear flow behavior at single fracture intersections.
637 *Journal of Hydrology*, 2016, 538(538):440-453.

638 Leung CTO, Zimmerman RW. Estimating the Hydraulic Conductivity of

639 Two-Dimensional Fracture Networks Using Network Geometric Properties. *Transport*
640 *in Porous Media*, 2012, 93(3):777-797.

641 Liu JS, Chen ZW, Elsworth D, Qu HY, Chen D. Interactions of multiple processes during
642 CBM extraction: A critical review. *International Journal of Coal Geology*, 2011,
643 87(3):175-189.

644 Liu RC, Li B, Jiang YJ. Critical hydraulic gradient for nonlinear flow through rock
645 fracture networks: The roles of aperture, surface roughness, and number of
646 intersections. *Advances in Water Resources*, 2016a, 88:53-65.

647 Liu RC, Li B, Jiang YJ, Huang N. Review: Mathematical expressions for estimating
648 equivalent permeability of rock fracture networks. *Hydrogeology Journal*, 2016c,
649 24(7):1-27.

650 Liu RC, Yu LY, Jiang YJ. Quantitative Estimates of Normalized Transmissivity and the
651 Onset of Nonlinear Fluid Flow Through Rough Rock Fractures. *Rock Mechanics &*
652 *Rock Engineering*, 2016b, 50(4):1-9.

653 Moutsopoulos KN. Exact and approximate analytical solutions for unsteady fully
654 developed turbulent flow in porous media and fractures for time dependent boundary
655 conditions. *Journal of Hydrology*, 2009, 369(1):78-89.

656 Olsson R, Barton N. An improved model for hydromechanical coupling during shearing
657 of rock joints. *International Journal of Rock Mechanics & Mining Sciences*, 2001,
658 38(3):317-329.

659 Oron A P, Berkowitz B. Flow in rock fractures: The local cubic law assumption
660 reexamined. *Water Resources Research*, 1998, 34(11): 2811-2825.

661 Qian J, Zhan H, Chen Z, Ye H. Experimental study of solute transport under non-Darcian
662 flow in a single fracture. *Journal of Hydrology*, 2011, 399(3):246-254.

663 Ranjith PG, Darlington W. Nonlinear single - phase flow in real rock joints. *Water*
664 *Resources Research*, 2007, 43(9):146-156.

665 Rong G, Hou D, Yang J, Cheng L, Zhou CB. Experimental study of flow characteristics in
666 non-mated rock fractures considering 3D definition of fracture surfaces. *Engineering*
667 *Geology*, 2017, 220:152-163.

668 Rong G, Yang J, Cheng L, Zhou CB. Laboratory investigation of nonlinear flow
669 characteristics in rough fractures during shear process. *Journal of Hydrology*, 2016,
670 541:1385-1394.

671 Schrauf TW, Evans DD. Laboratory studies of gas flow through a single natural fracture.
672 *Water Resources Research*, 1986, 22(7):1038-1050.

673 Tse R, Cruden D M. Estimating joint roughness coefficients. *International Journal of*
674 *Rock Mechanics & Mining Sciences & Geomechanics Abstracts*, 1979, 16(5):303-307.

675 Tzelepis V, Moutsopoulos K N, Papaspyros J N E, Tsihrintzis V A. Experimental
676 investigation of flow behavior in smooth and rough artificial fractures. *Journal of*
677 *Hydrology*, 2015, 521(2):108-118.

678 Wang CG, Liu JS, Feng JL, Wei MY, Wang CC, Jiang YJ. Effects of Gas Diffusion from
679 Fractures to Coal Matrix on the Evolution of Coal Strains: Experimental Observations.
680 *International Journal of Coal Geology*, 2016, 162:74-84.

681 Wang CG, Zhai PC, Chen ZW, Liu JS, Wang LS, Xie J. Experimental Study of Coal
682 Matrix-fracture Interaction under Constant Volume Boundary Condition. *International*
683 *Journal of Coal Geology*, 2017,181,124-132.

684 Wang L, Cardenas MB, Slotke DT, Ketcham RA, Sharp JM. Modification of the Local
685 Cubic Law of fracture flow for weak inertia, tortuosity, and roughness. *Water*
686 *Resources Research*, 2015, 51(4):2064-2080.

687 Wang M, Chen YF, Ma GW, Zhou JQ, Zhou CB. Influence of surface roughness on
688 nonlinear flow behaviors in 3D self-affine rough fractures: Lattice Boltzmann
689 simulations. *Advances in Water Resources*, 2016, 96:373-388.

690 Wei W, Xia Y. Geometrical, fractal and hydraulic properties of fractured reservoirs: A
691 mini-review. *Adv. Geo Energy*, 2017, 1(1): 31-38.

692 Wu Y, Liu JS, Chen ZW, Elsworth D, Pone D. A dual poroelastic model for
693 CO₂-enhanced coalbed methane recovery. *International Journal of Coal Geology*, 2011,
694 86(2-3): 177-189.

695 Xia CC, Qian X, Lin P, Xiao WM, Gui Y. Experimental Investigation of Nonlinear Flow
696 Characteristics of Real Rock Joints under Different Contact Conditions. *Journal of*
697 *Hydraulic Engineering*, 2016, 143(3).

698 Xiong XB, Li B, Jiang YJ, Koyama T, Zhang CH. Experimental and numerical study of
699 the geometrical and hydraulic characteristics of a single rock fracture during shear.
700 *International Journal of Rock Mechanics & Mining Sciences*, 2011, 48(8):1292-1302.

701 Yeo I W, De Freitas, M H, Zimmerman R W. Effect of shear displacement on the
702 aperture and permeability of a rock fracture. *International Journal of Rock Mechanics*
703 *& Mining Sciences*, 1998, 35(8), 1051–1070.

704 Yin Q, Jing HW, Liu RC, Ma GW, Yu LY, Su HJ. Experimental Study on
705 Stress-Dependent Nonlinear Flow Behavior and Normalized Transmissivity of Real
706 Rock Fracture Networks. *Geofluids*, 2018.

707 Yin Q, Ma GW, Jing HW, Wang HD, Su HJ, Wang YC, Liu RC. Hydraulic properties of
708 3D rough-walled fractures during shearing: an experimental study. *Journal of*
709 *Hydrology*, 2017, 555.

710 Yu LY, Liu RC, Jiang YJ. A Review of Critical Conditions for the Onset of Nonlinear

711 Fluid Flow in Rock Fractures. *Geofluids*, 2017.

712 Zeng Z, Grigg R. A Criterion for Non-Darcy Flow in Porous Media. *Transport in Porous*
713 *Media*, 2006, 63(1):57-69.

714 Zhang W, Dai B, Liu Z, Zhou C. A pore-scale numerical model for non-Darcy fluid flow
715 through rough-walled fractures. *Computers & Geotechnics*, 2017, 87:139-148.

716 Zhang ZY, Nemeik J. Fluid flow regimes and nonlinear flow characteristics in deformable
717 rock fractures. *Journal of Hydrology*, 2013, 477(1):139-151.

718 Zhou JQ, Hu SH, Fang S, Chen YF, Zhou CB. Nonlinear flow behavior at low Reynolds
719 numbers through rough-walled fractures subjected to normal compressive loading.
720 *International Journal of Rock Mechanics & Mining Sciences*. 2015, 80:202–218.

721 Zimmerman RW, Al-Yaarubi A, Pain CC, Grattoni CA. Non-linear regimes of fluid flow
722 in rock fractures. *International Journal of Rock Mechanics & Mining Sciences*, 2004,
723 41(3):384.

724 Zimmerman RW, Bodvarsson G S. Hydraulic conductivity of rock fractures. *Transport in*
725 *Porous Media*, 1996, 23(1):1-30.

726

Table 1 Physico-mechanical properties of plaster and acrylic specimen

Physico-mechanical properties	Index	Unit	Plaster specimen	Acrylic specimen
Density	ρ	g/cm ³	2.066	1.192
Compressive strength	σ_c	MPa	38.5	91
Modulus of elasticity	E_s	MPa	28700	24744
Poisson's ratio	ν	–	0.23	0.19
Tensile strength	σ_t	MPa	2.5	5.6
Cohesion	c	MPa	5.3	19.9
Internal friction angle	φ	°	60	42.8

Table 2 Experimental results of the characteristic mechanical parameter during shear flow tests

Specimen NO.	Case NO.	JRC	σ_n (MPa)	d_{peak} (mm)	τ_{peak} (MPa)	K_s (MPa/mm)	d_{normal} (mm)	d_{contra} (mm)
G1	G1-1	3.21	0.5	0.844	0.515	0.61	1.080	-0.063
	G1-2		1.0	1.167	1.003	0.86	0.850	-0.080
	G1-3		1.5	1.244	1.407	1.13	0.687	-0.118
	G1-4		2.0	1.417	1.789	1.26	0.561	-0.173
G2	G2-1	5.62	0.5	0.878	0.652	0.74	1.309	-0.003
	G2-2		1.0	1.034	1.136	1.10	1.235	-0.042
	G2-3		1.5	1.195	1.497	1.25	1.09	-0.063
	G2-4		2.0	1.228	1.834	1.49	0.919	-0.073
G3	G3-1	7.36	0.5	0.94	0.770	0.82	1.835	-0.007
	G3-2		1.0	1.059	1.204	1.14	1.212	-0.032
	G3-3		1.5	1.098	1.536	1.40	0.999	-0.060
	G3-4		2.0	1.141	2.070	1.80	0.956	-0.143
G4	G4-1	12.16	0.5	0.892	0.813	0.91	2.472	-0.013
	G4-2		1.0	1.074	1.485	1.40	2.200	-0.064
	G4-3		1.5	1.152	1.872	1.63	1.685	-0.074
	G4-4		2.0	1.168	2.141	1.83	1.269	-0.073

Table 3 Measured results of J_c , Re_c and β for different roughness fractures during shear

Specime N0.	σ_n (MPa)	d (mm)	J_c	Re_c	β	Specime N0.	σ_n (MPa)	d (mm)	J_c	Re_c	β
G1-1	0.5	1	1.154	1.83	0.06051	G3-1	0.5	1	0.496	1.40	0.07916
		3	4.481	21.16	0.00520			3	2.327	15.74	0.00706
		5	5.798	30.11	0.00367			5	2.573	24.37	0.00456
		7	6.567	40.40	0.00275			7	2.894	29.16	0.00381
		9	6.934	48.73	0.00227			9	3.048	39.12	0.00284
G1-2	1.0	1	1.208	1.52	0.07298	G3-2	1.0	1	0.736	1.19	0.09322
		3	6.701	19.56	0.00568			3	2.859	13.27	0.00837
		5	7.322	28.13	0.00395			5	3.627	19.87	0.00559
		7	7.764	35.05	0.00317			7	3.989	24.00	0.00463
		9	8.187	44.44	0.00250			9	4.379	30.28	0.00367
G1-3	1.5	1	1.483	1.65	0.06710	G3-3	1.5	1	0.957	1.24	0.08930
		3	7.331	17.44	0.00637			3	3.339	11.78	0.00943
		5	7.733	25.90	0.00429			5	3.836	17.31	0.00642
		7	8.182	30.69	0.00362			7	4.499	21.57	0.00515
		9	8.493	38.85	0.00286			9	4.659	29.55	0.00376
G1-4	2.0	1	1.501	1.52	0.07311	G3-4	2.0	1	1.216	1.25	0.08909
		3	8.578	16.53	0.00672			3	3.926	10.17	0.01092
		5	8.712	24.91	0.00446			5	4.296	12.93	0.00859
		7	9.335	28.93	0.00384			7	4.978	18.24	0.00609
		9	9.592	38.75	0.00274			9	5.073	26.27	0.00423
G2-1	0.5	1	0.979	1.56	0.07117	G4-1	0.5	1	0.289	1.27	0.08733
		3	3.417	16.20	0.00686			3	0.982	12.23	0.00908
		5	3.873	21.45	0.00518			5	1.126	18.89	0.00588
		7	4.113	27.57	0.00403			7	1.266	24.47	0.00454
		9	4.356	35.27	0.00315			9	1.312	29.71	0.00374
G2-2	1.0	1	1.156	1.48	0.07520	G4-2	1.0	1	0.448	1.34	0.08287
		3	3.579	13.65	0.00814			3	1.010	9.33	0.01191
		5	4.276	23.89	0.00465			5	1.352	16.93	0.00656
		7	4.490	27.23	0.00408			7	1.588	23.15	0.0048
		9	4.698	32.21	0.00345			9	1.651	27.3	0.00407
G2-3	1.5	1	1.257	1.35	0.08235	G4-3	1.5	1	0.580	1.28	0.08711
		3	4.286	13.45	0.00826			3	1.194	8.10	0.01371
		5	4.728	21.66	0.00513			5	1.599	14.81	0.0075

		7	4.828	25.37	0.00438			7	1.805	21.83	0.00509
		9	5.006	29.71	0.00374			9	1.959	24.05	0.00462
G2-4	2.0	1	1.312	1.27	0.08717	G4-4	2.0	1	0.846	1.15	0.09651
		3	4.796	12.55	0.00885			3	1.500	8.24	0.01347
		5	5.086	20.39	0.00545			5	1.735	13.4	0.00829
		7	5.211	24.86	0.00447			7	1.930	19.7	0.00564
		9	5.391	28.06	0.00396			9	2.028	23.1	0.00481

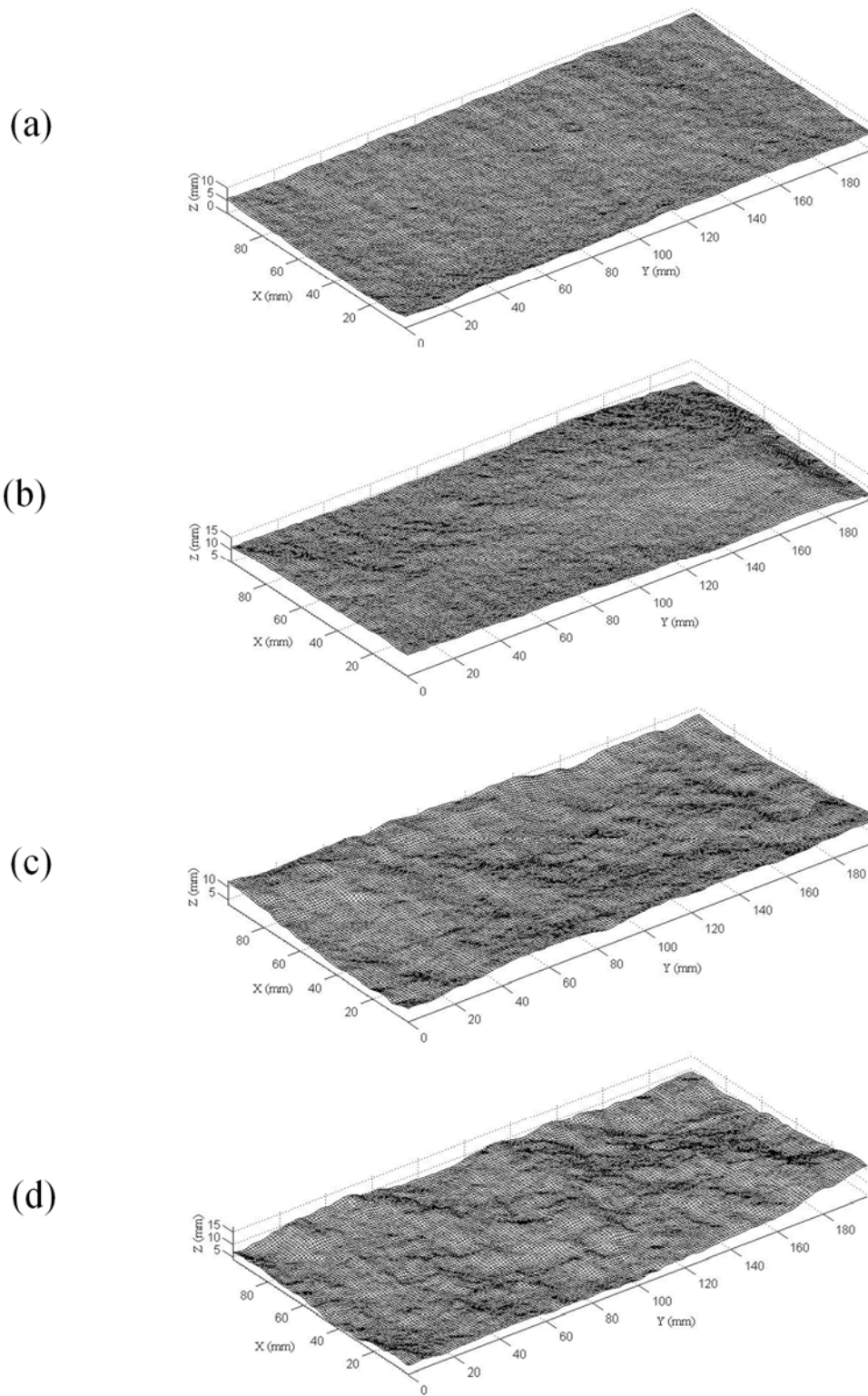


Fig. 1 Scanning graphs of four fracture surface: (a) Specimen G1, $JRC = 3.21$; (b) specimen G2, $JRC = 5.62$; (c) specimen G3, $JRC = 7.36$; and (d) specimen G4, $JRC = 12.16$.

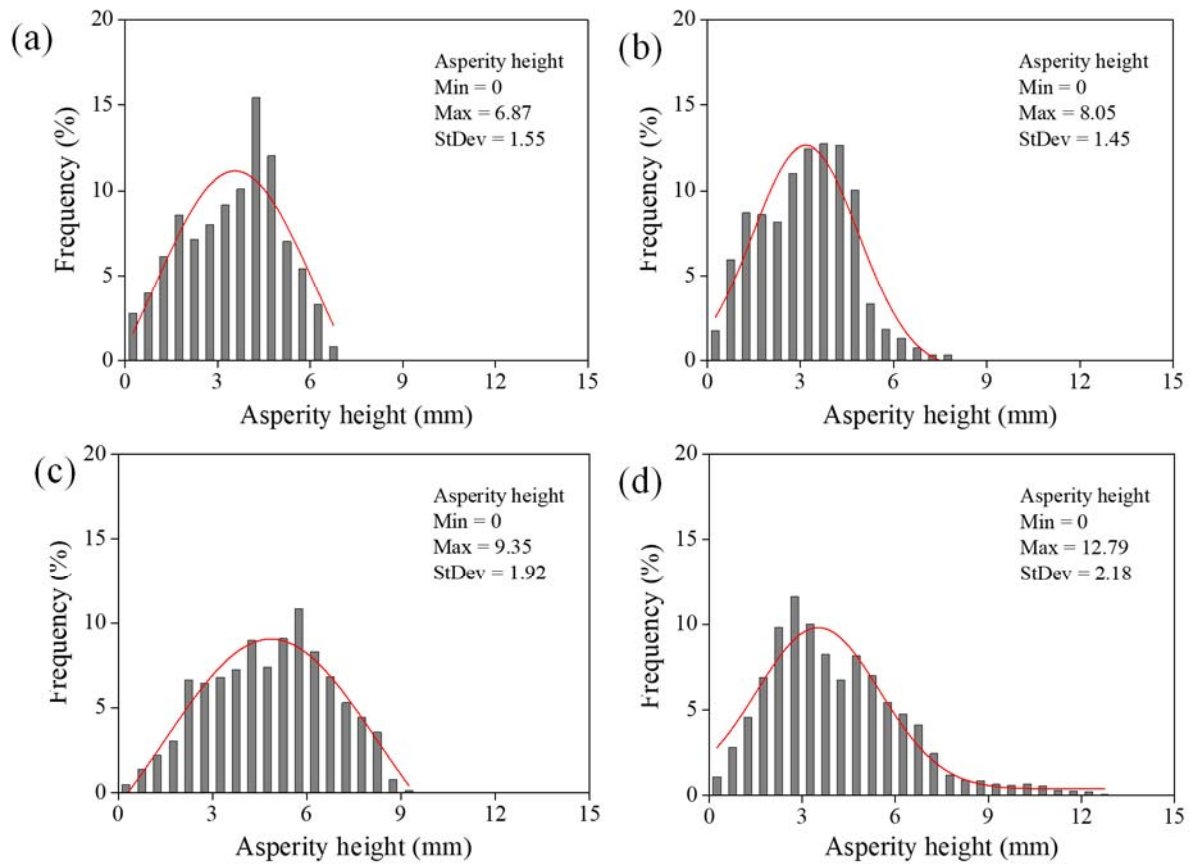


Fig. 2 Frequency versus asperity height of four fracture surfaces: (a) for fracture G1; (b) for fracture G2; (c) for fracture G3; and (d) for fracture G4.

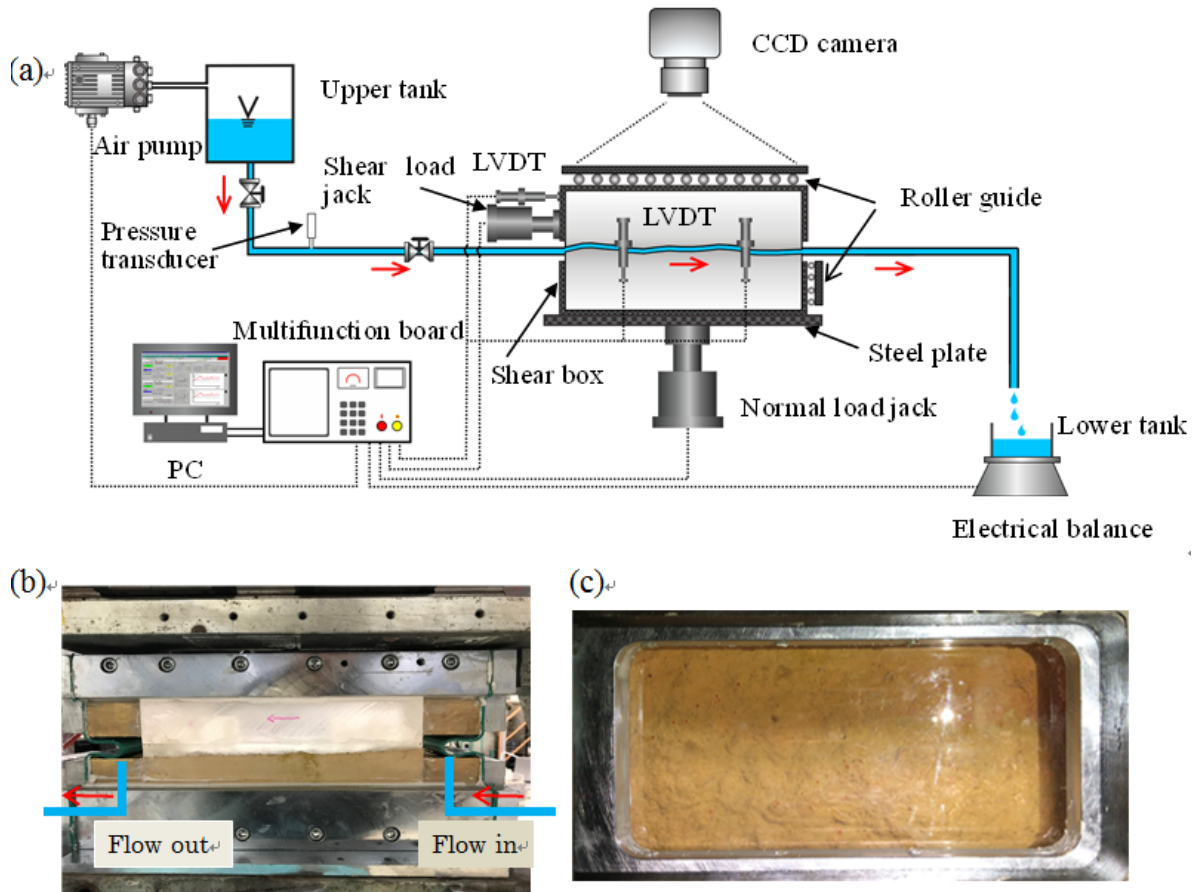


Fig. 3 Schematic view of the coupled shear-flow test apparatus (arrow represents water flow direction). (a) Hydraulic testing system; (b) Side view of the shear box; and (c) Top view of shear box.

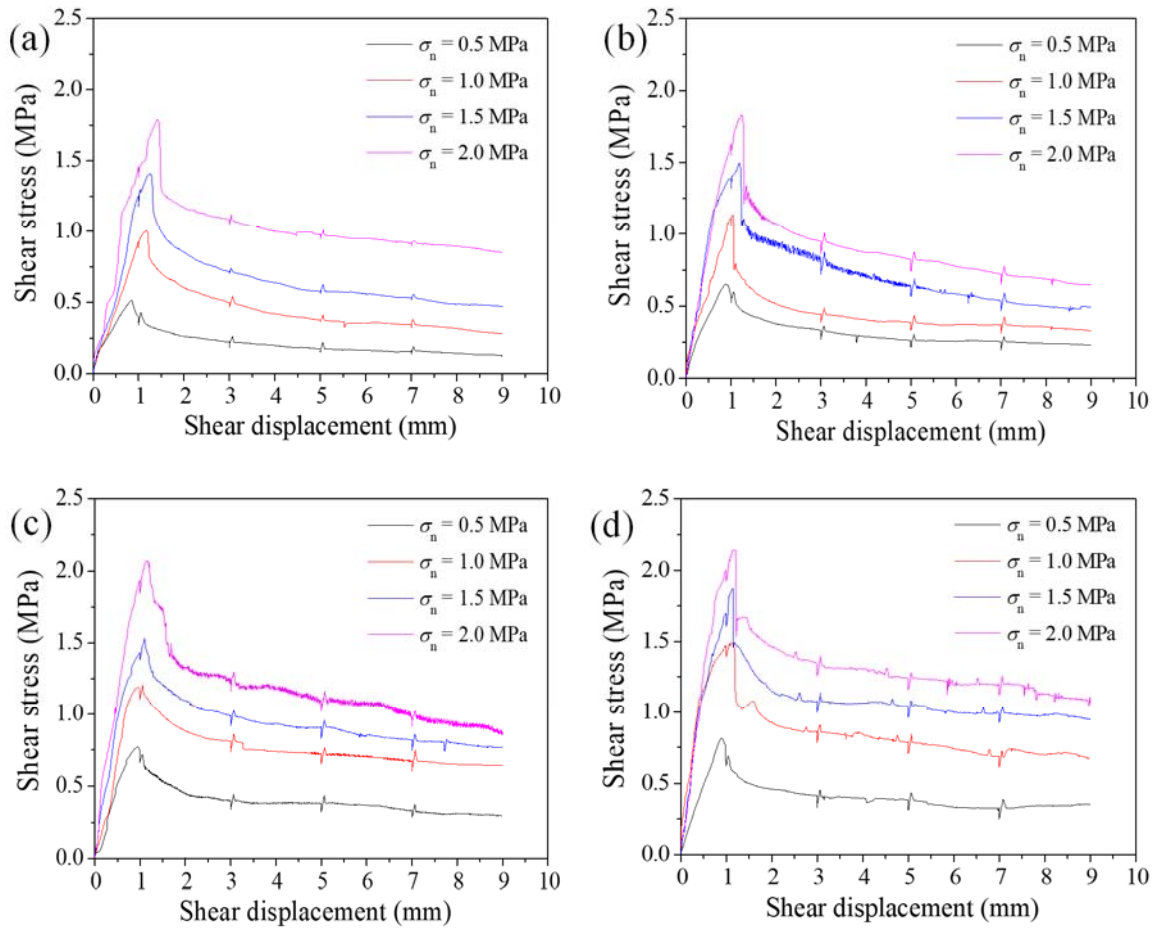


Fig. 4 Shear stress versus shear displacement of fracture specimens under different normal stresses: (a) for fracture G1; (b) for fracture G2; (c) for fracture G3 and (d) for fracture G4.

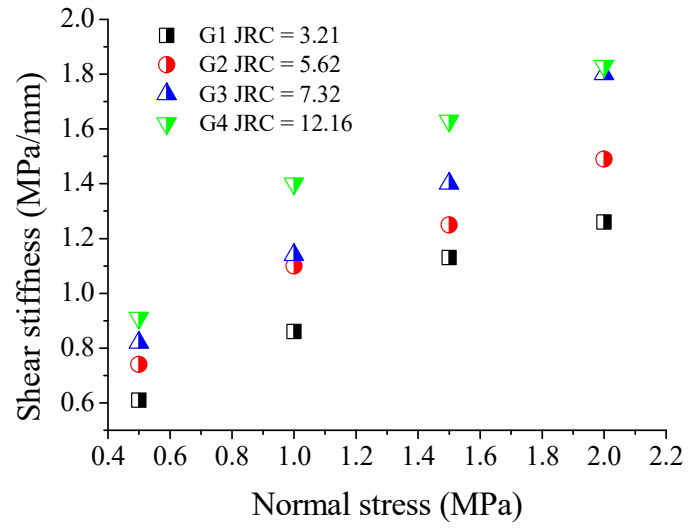


Fig. 5 Relationship between shear stiffness and normal stress of different fracture specimens

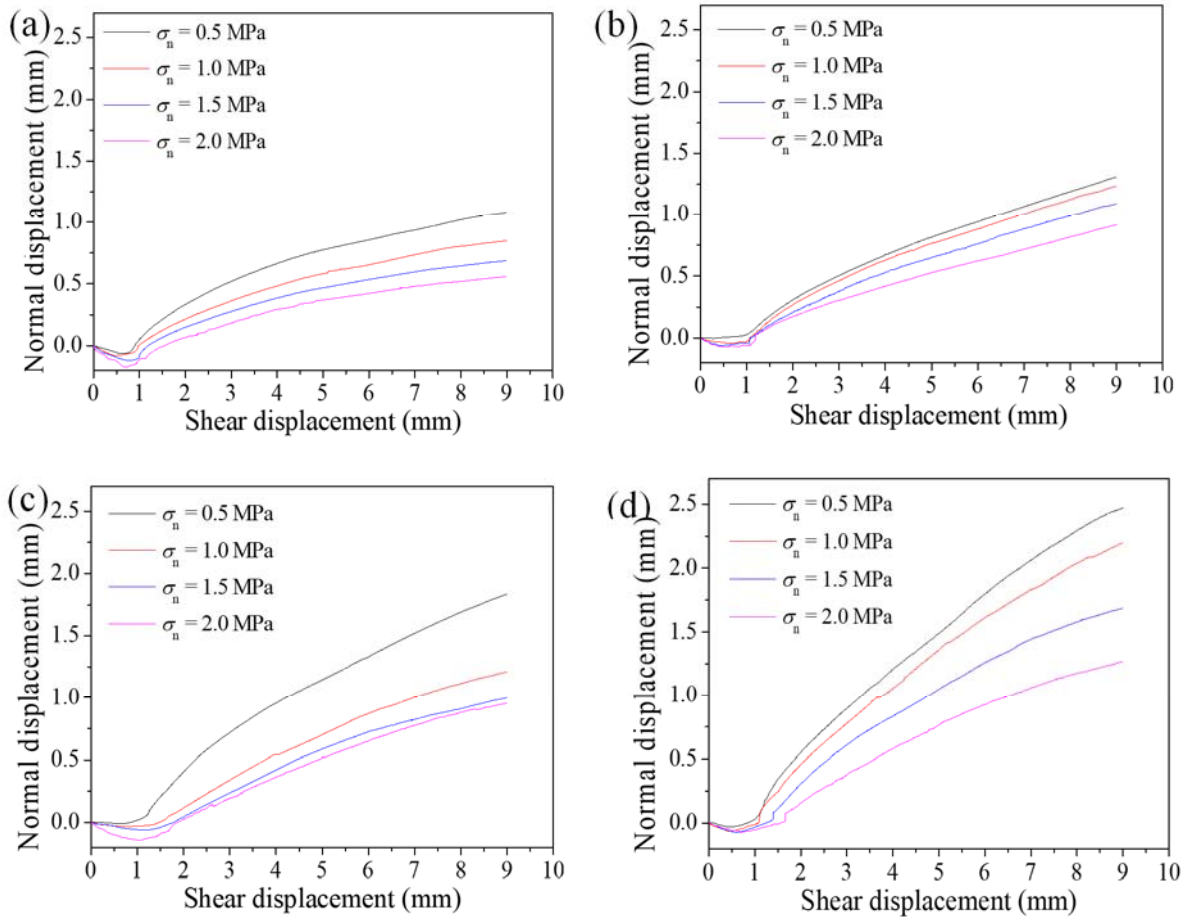
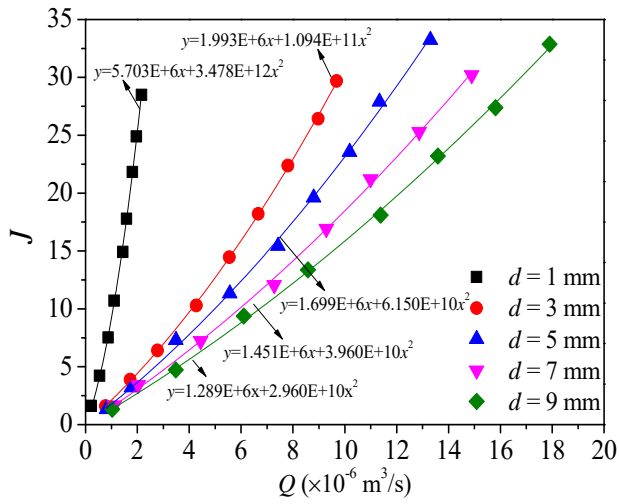
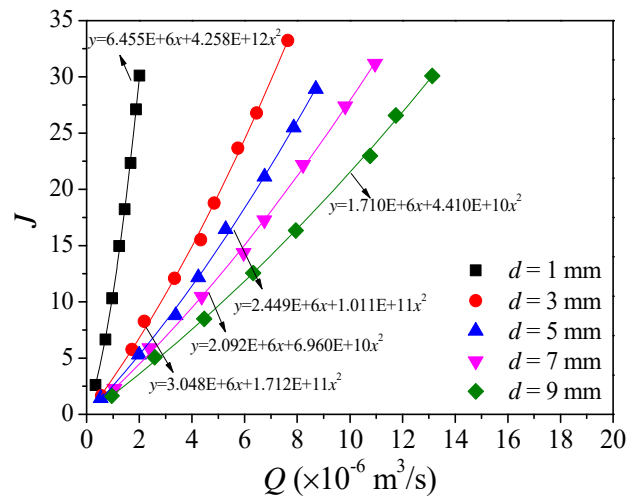


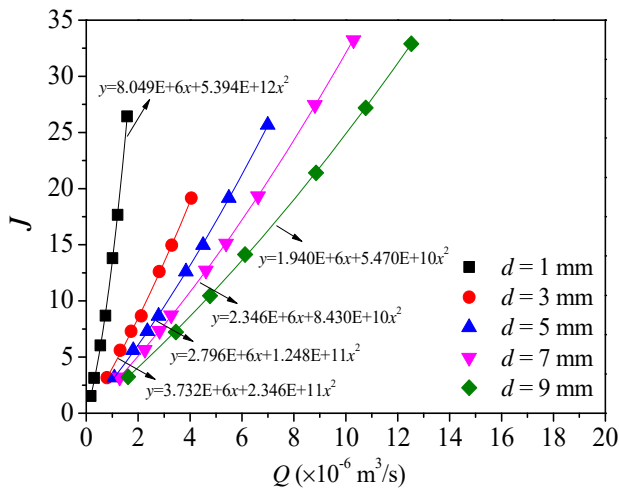
Fig. 6 Normal displacement versus shear displacement of fracture specimens under different normal stresses: (a) for fracture G1; (b) for fracture G2; (c) for fracture G3 and (d) for fracture G4.



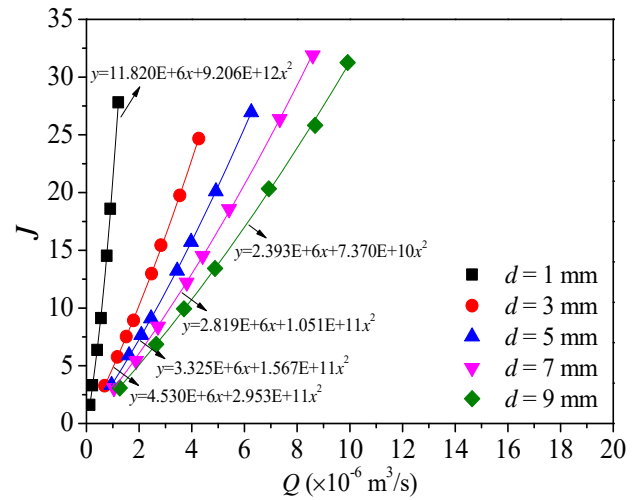
(a) $\sigma_n = 0.5$ MPa



(b) $\sigma_n = 1.0$ MPa



(c) $\sigma_n = 1.5$ MPa



(d) $\sigma_n = 2.0$ MPa

Fig. 7 Relationships between hydraulic gradient (J) and volumetric flow rate (Q) for G1 under different normal stresses.

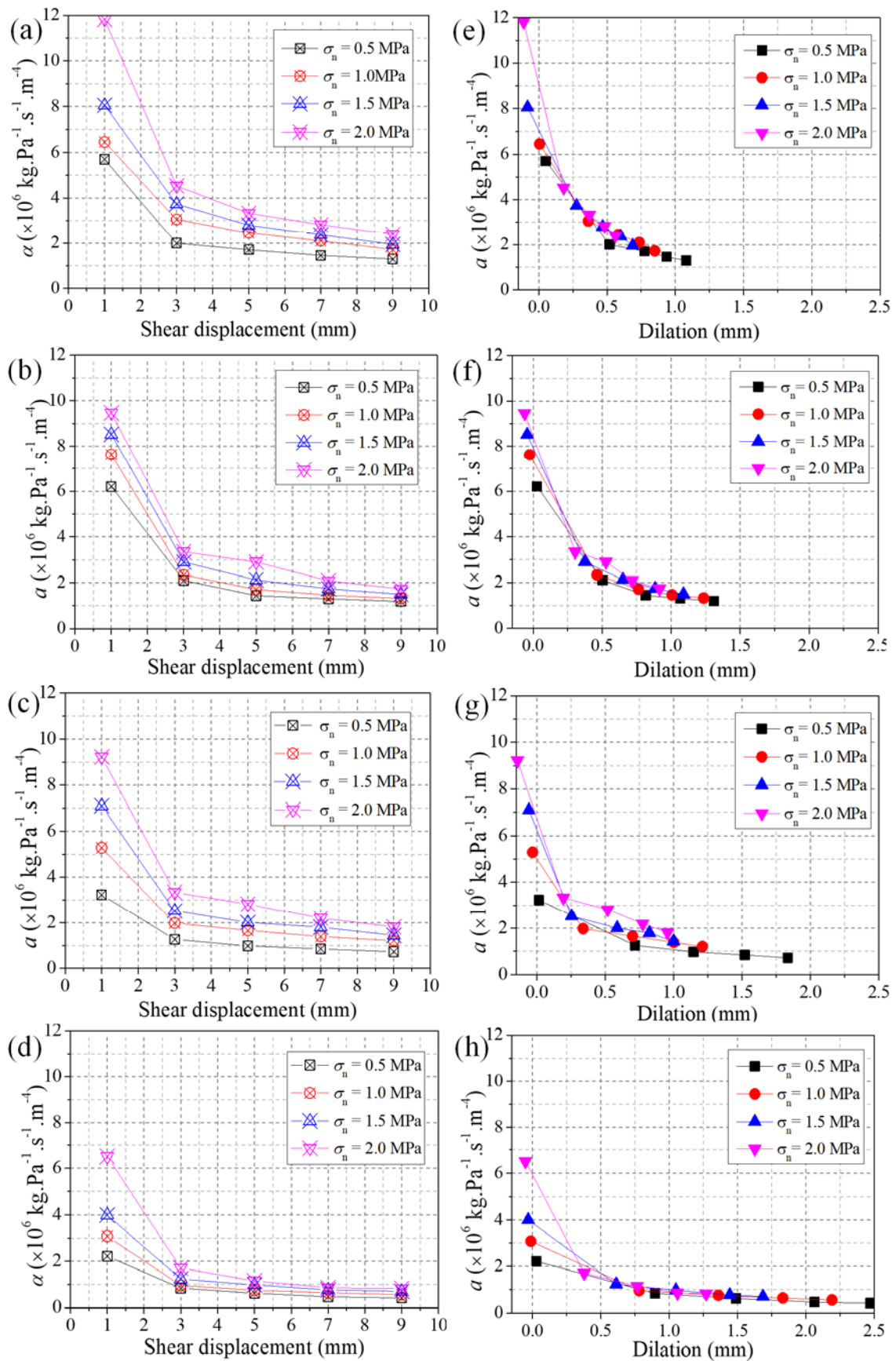


Fig. 8 Relationships between linear coefficient a and shear displacement: (a) G1, (b) G2, (c) G3, and (d) G4. Relationships between linear coefficient a and dilation: (e) G1, (f) G2, (g) G3, and (h) G4.

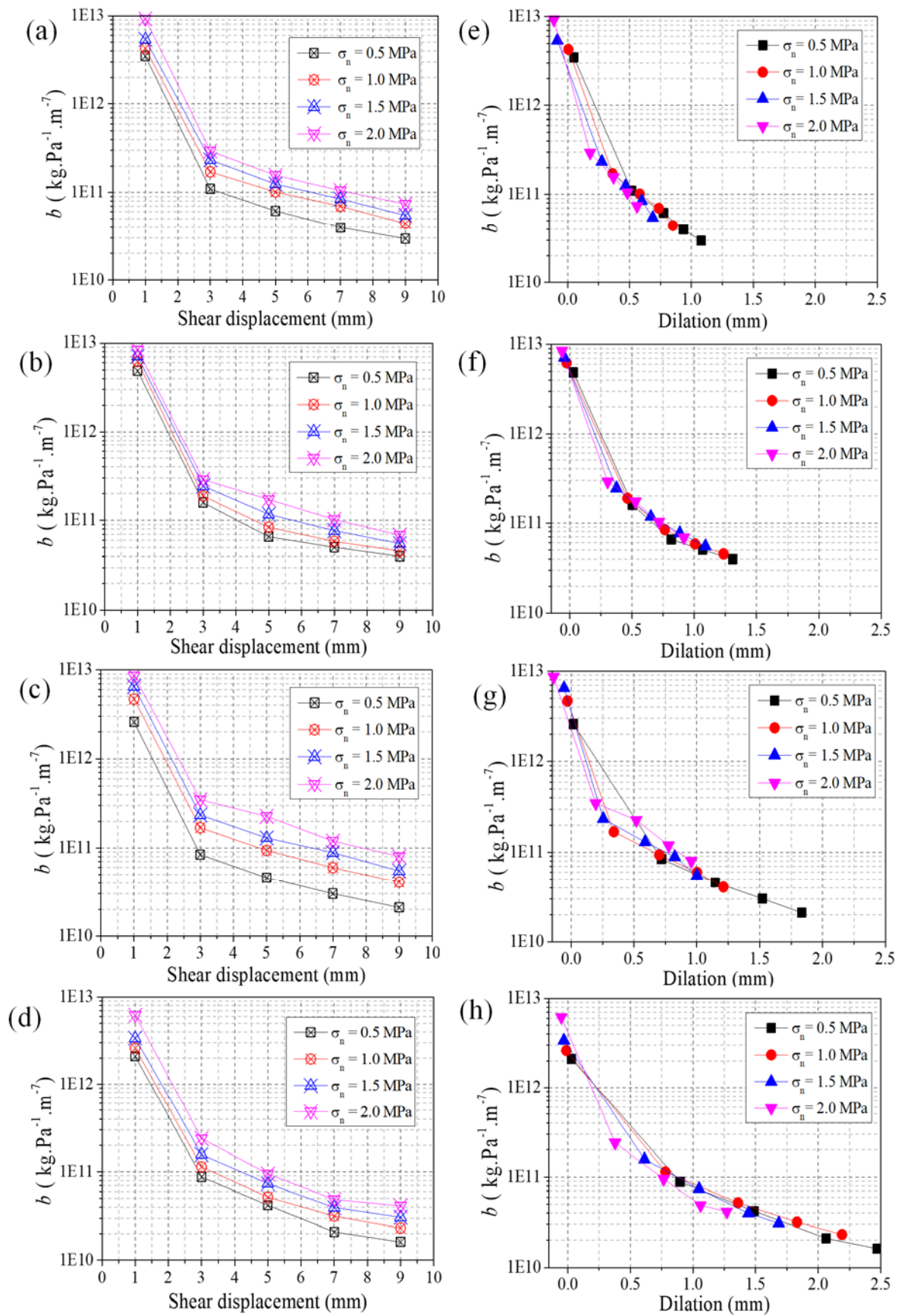


Fig. 9 Relationships between linear coefficient b and shear displacement: (a) G1, (b) G2, (c) G3, and (d) G4. Relationships between linear coefficient b and dilation: (e) G1, (f) G2, (g) G3, and (h) G4.

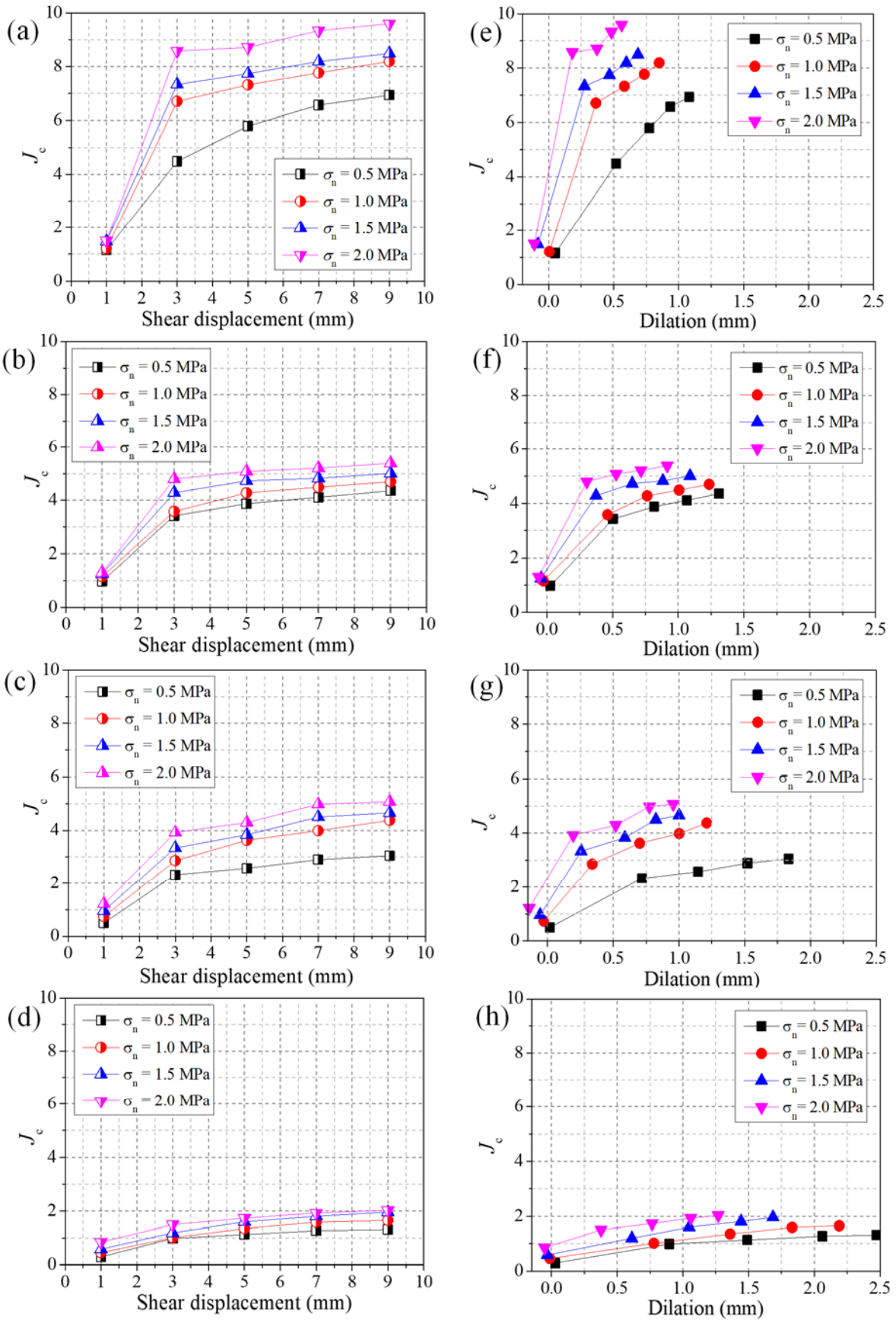


Fig. 10 Relationships between critical hydraulic gradient J_c and shear displacement d : (a) G1, (b) G2, (c) G3, and (d) G4. Relationships between critical hydraulic gradient J_c and dilation: (e) G1, (f) G2, (g) G3, and (h) G4.

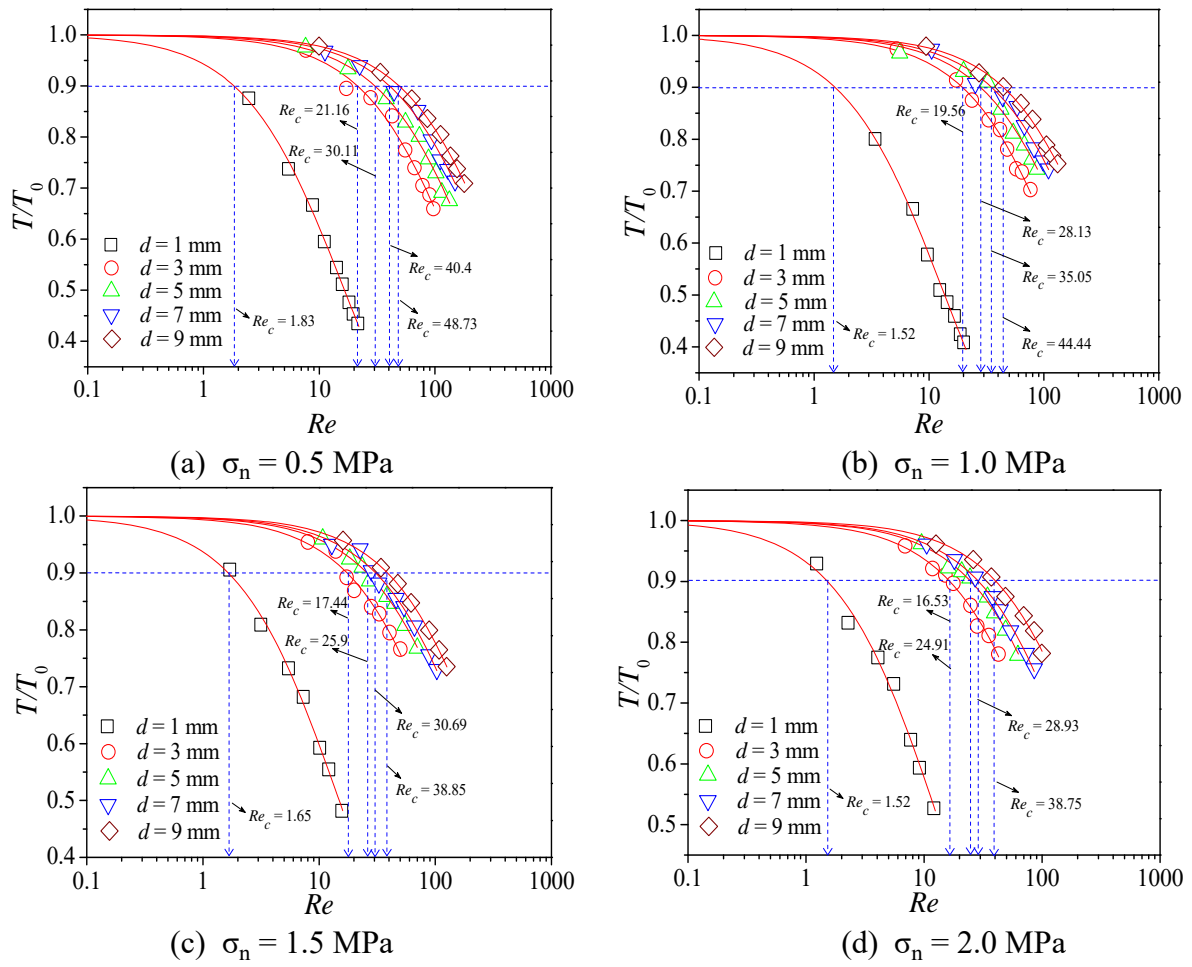


Fig. 11 Relationships between normalized transmissivity (T/T_0) and Reynolds number (Re) for G1 under different normal stresses.

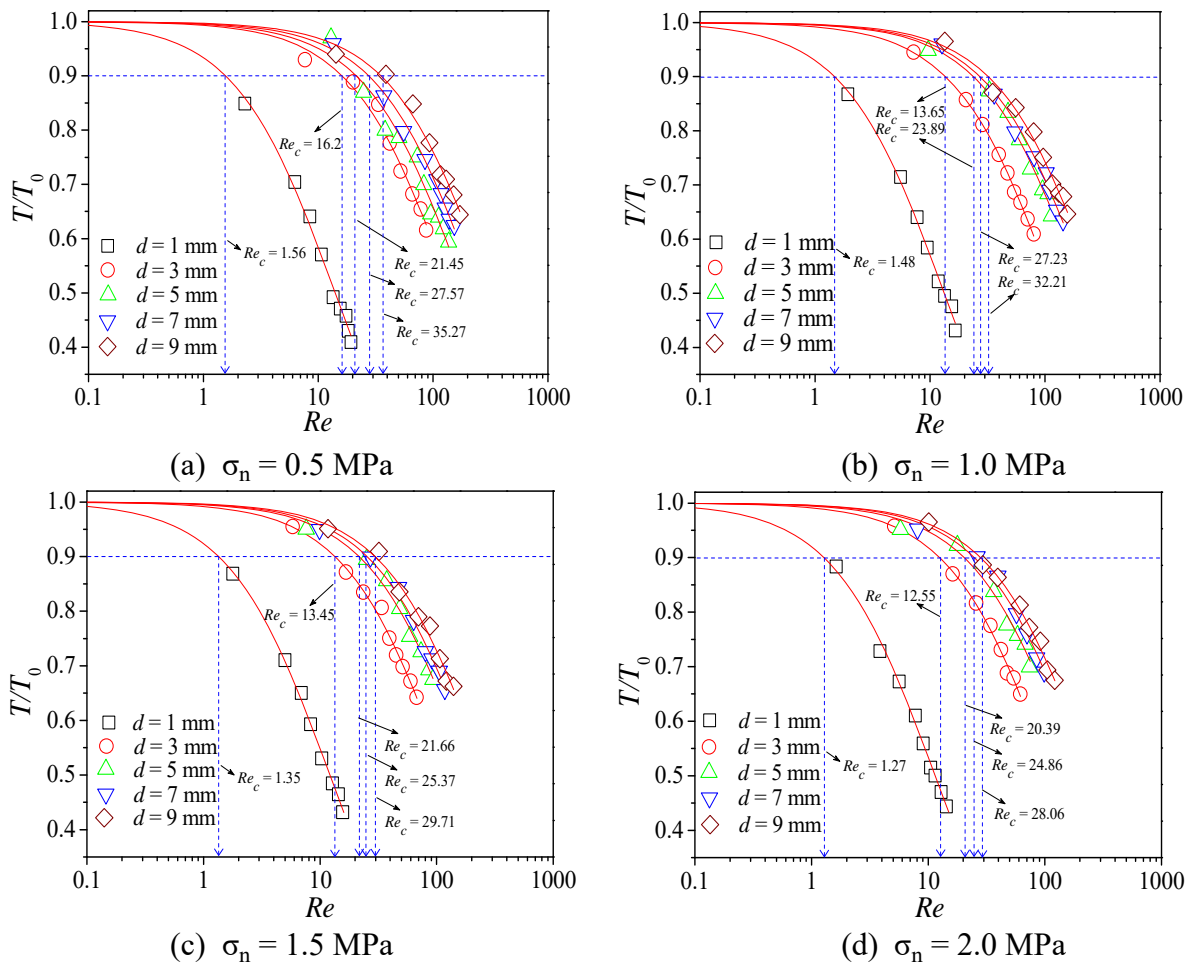


Fig. 12 Relationships between normalized transmissivity (T/T_0) and Reynolds number (Re) for G2 under different normal stresses.

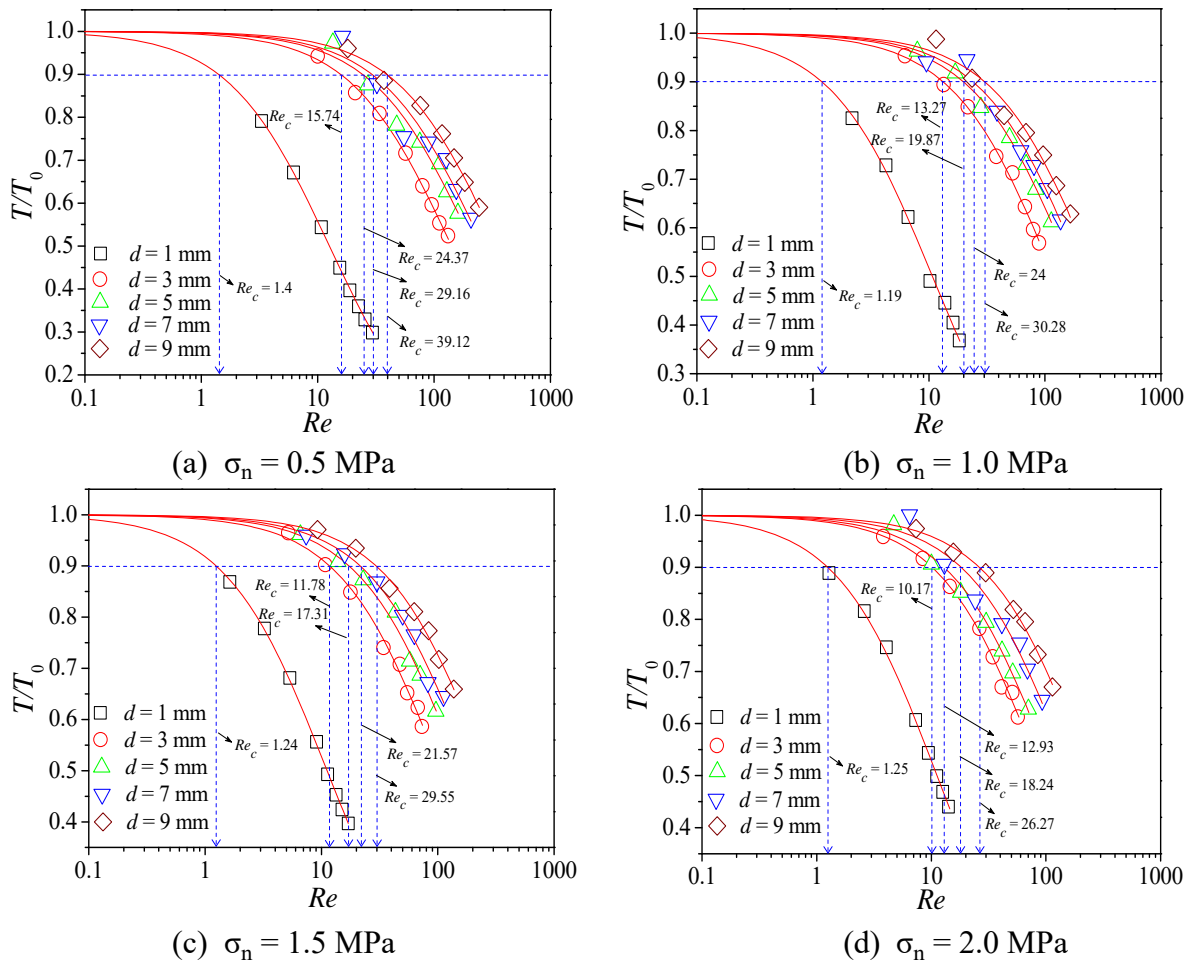


Fig. 13 Relationships between normalized transmissivity (T/T_0) and Reynolds number (Re) for G3 under different normal stresses.

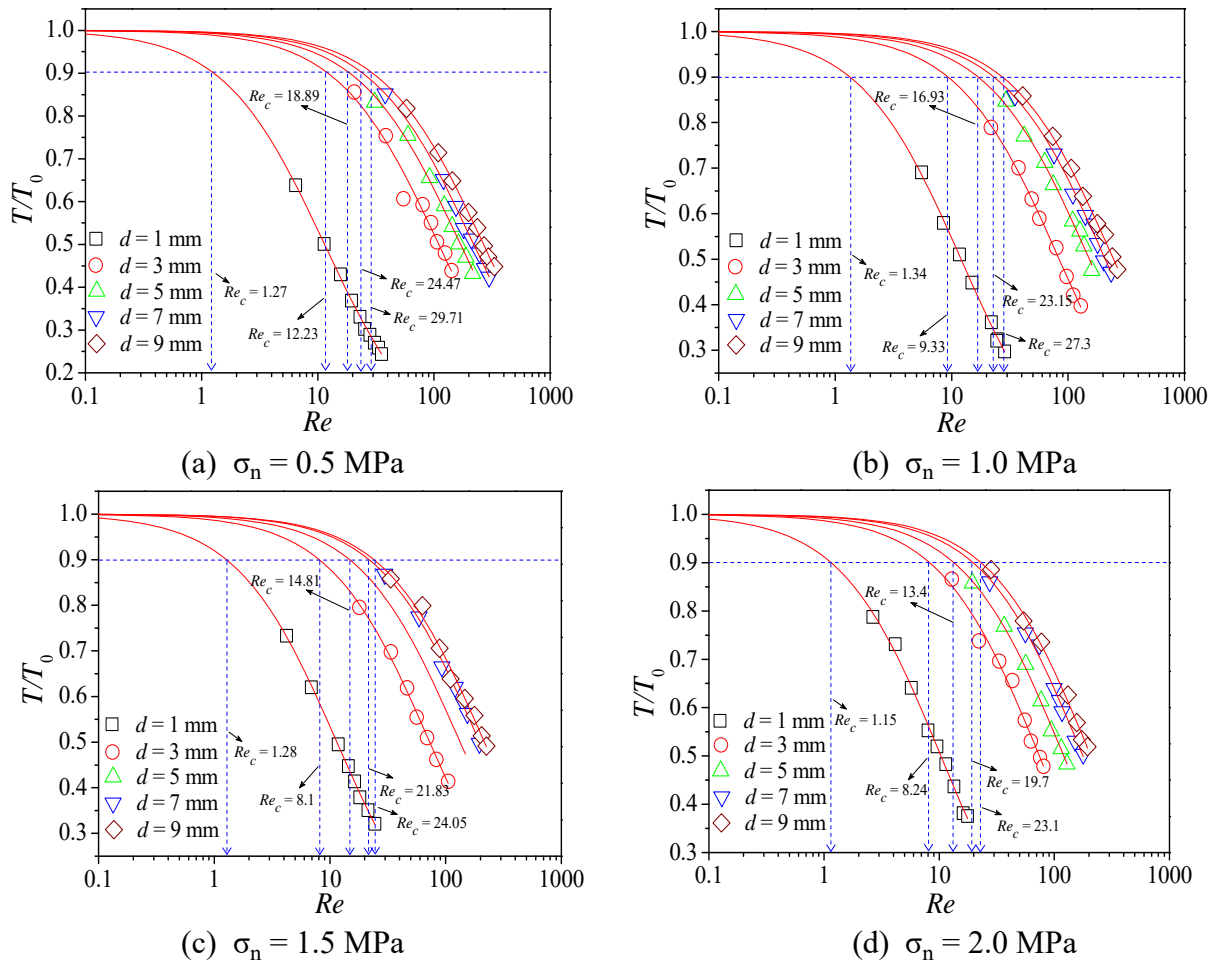


Fig. 14 Relationships between normalized transmissivity (T/T_0) and Reynolds number (Re) for G4 under different normal stresses.

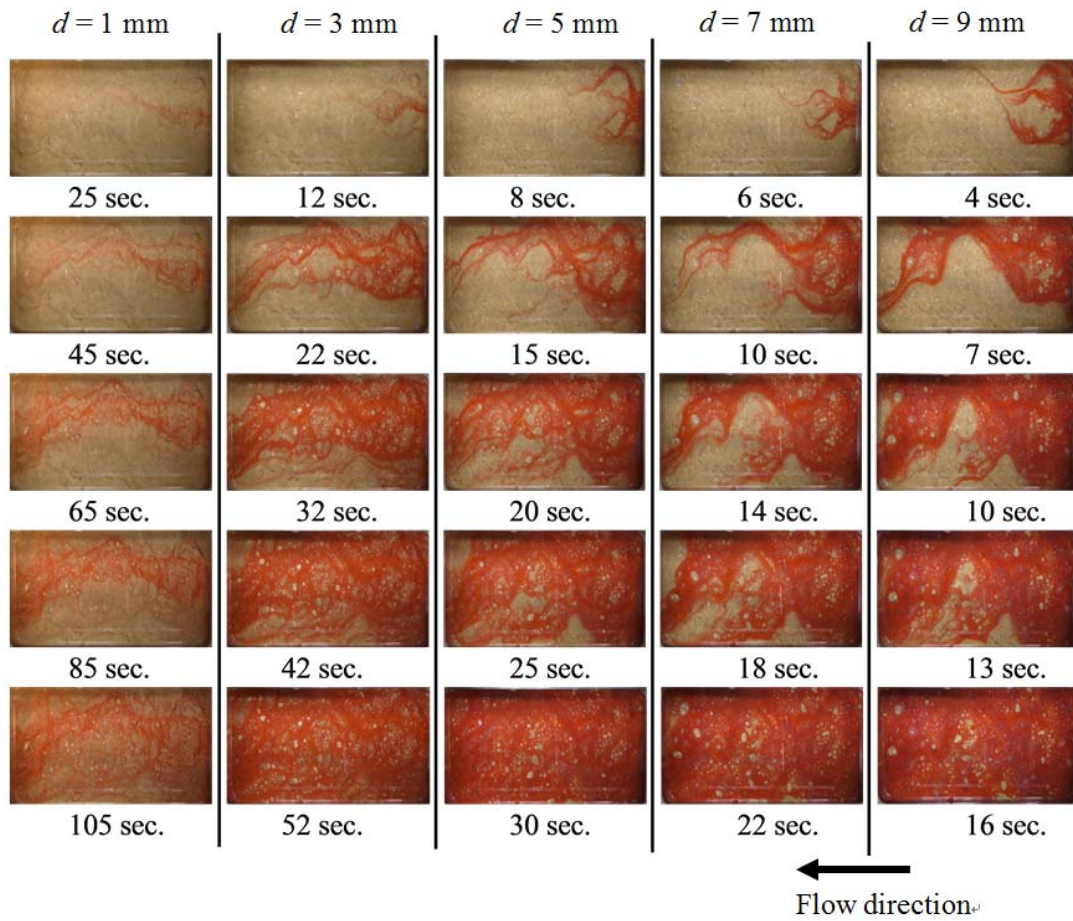


Fig. 15 Visualizations of the process of fluid flow in specimen G3-1 at different shear displacement.

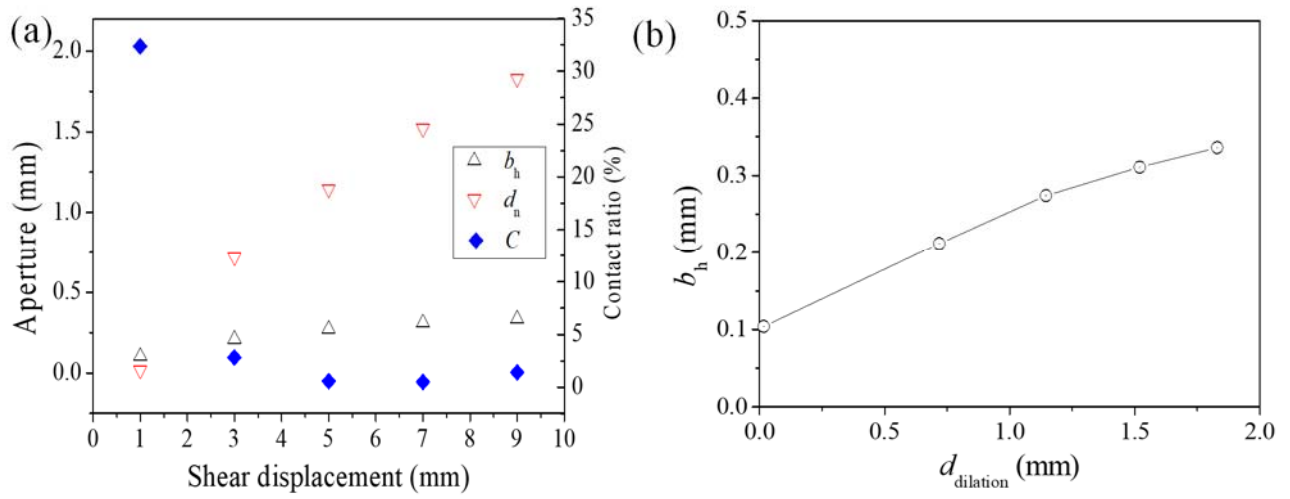
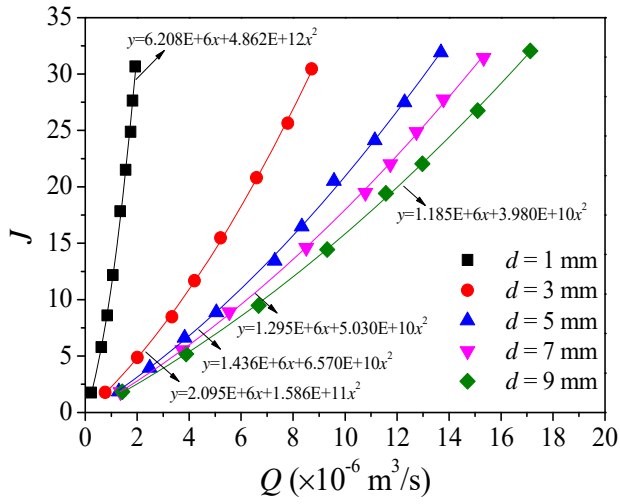
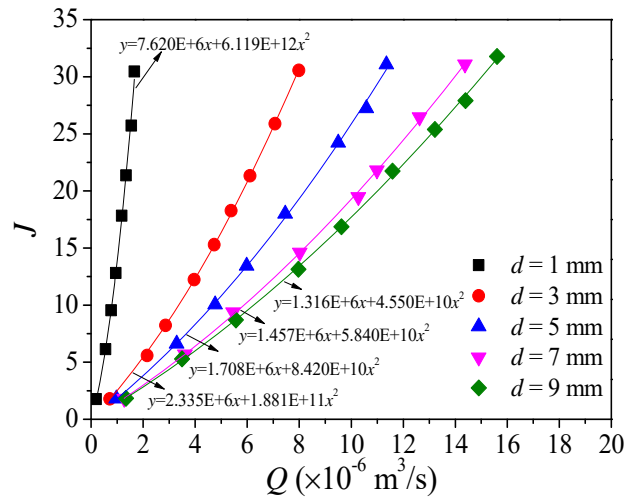


Fig. 16 (a) Evolution of hydraulic aperture b_h , dilation $d_{dilation}$ and contact ratio C of the specimen G3-1 during shear. (b) Relationship between hydraulic aperture b_h and dilation $d_{dilation}$.

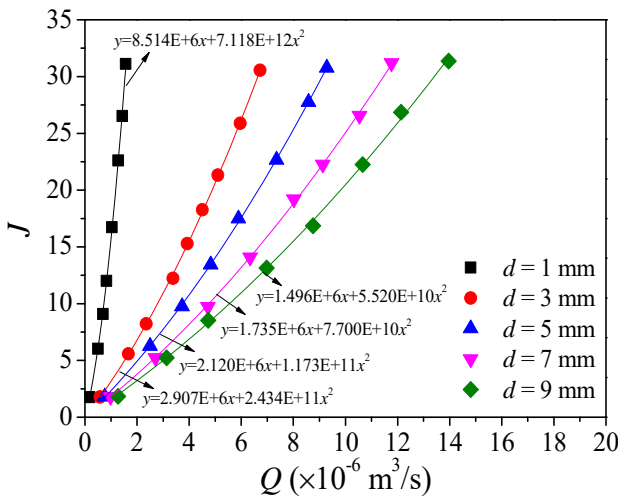
Appendix A



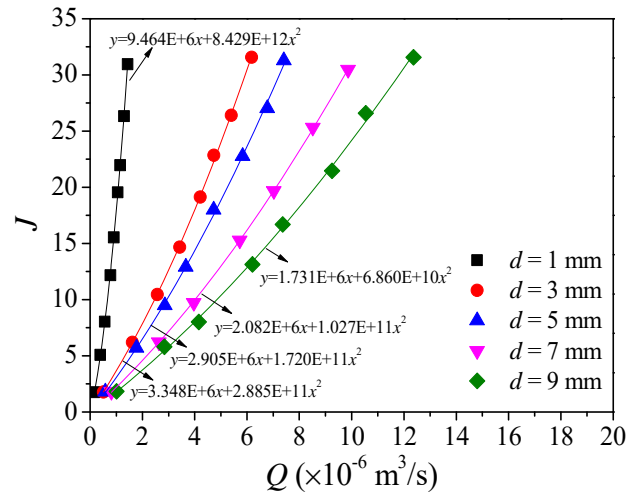
(a) G2 $\sigma_n = 0.5$ MPa



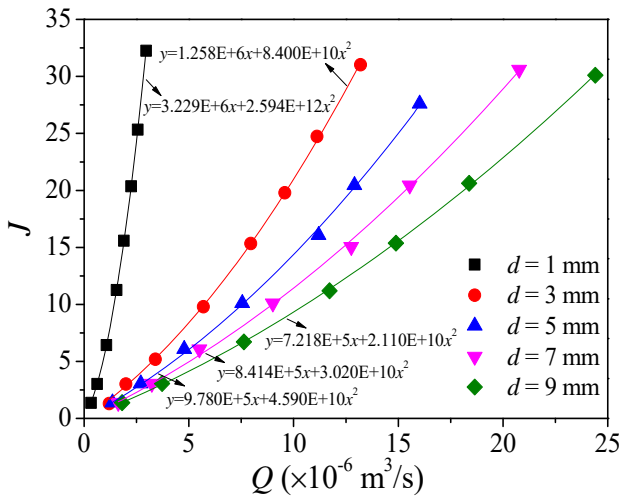
(b) G2 $\sigma_n = 1.0$ MPa



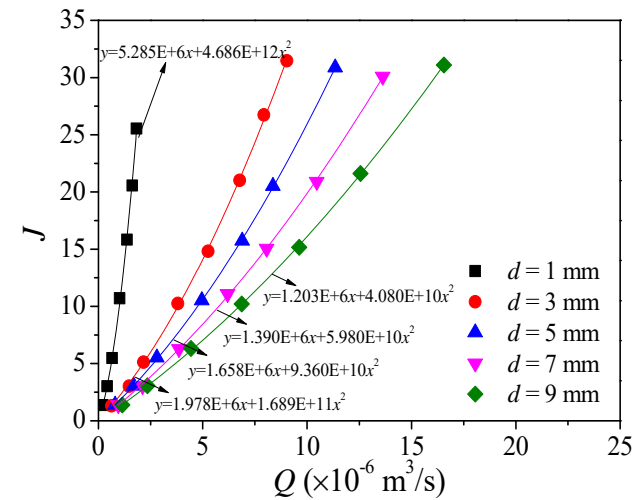
(c) G2 $\sigma_n = 1.5$ MPa



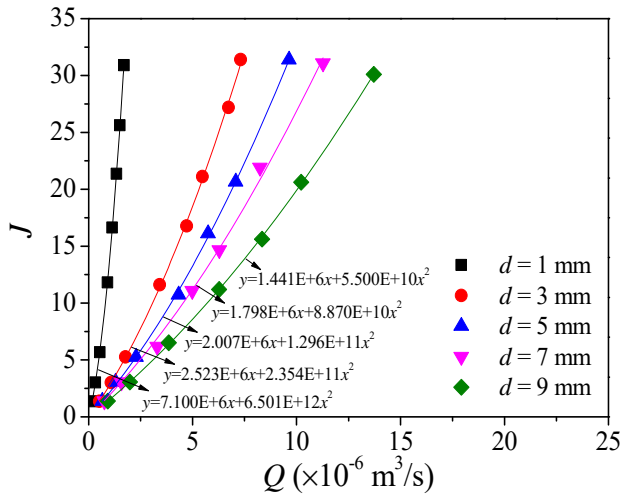
(d) G2 $\sigma_n = 2.0$ MPa



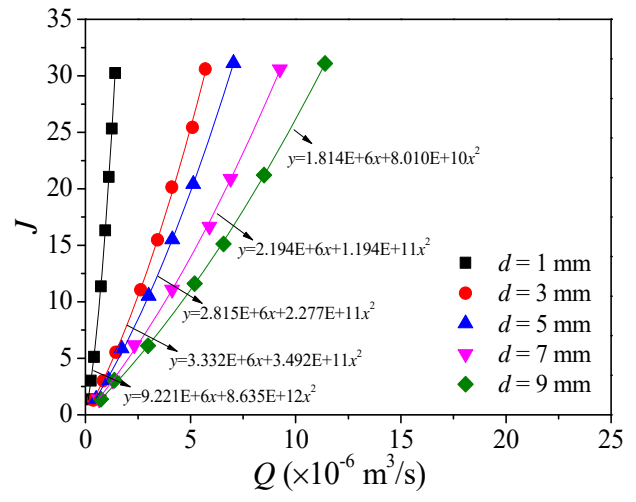
(e) G3 $\sigma_n = 0.5$ MPa



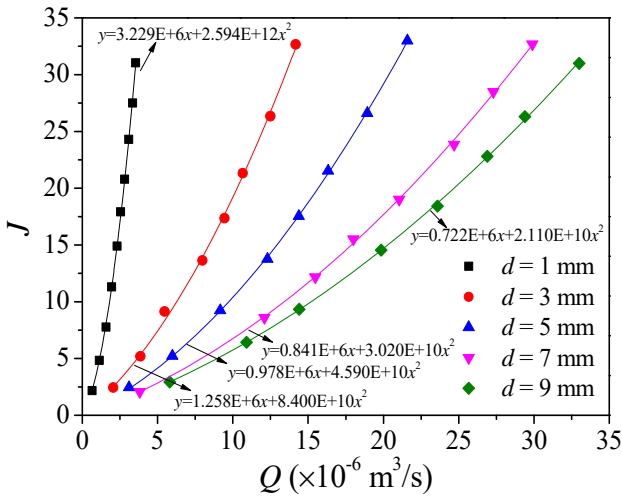
(f) G3 $\sigma_n = 1.0$ MPa



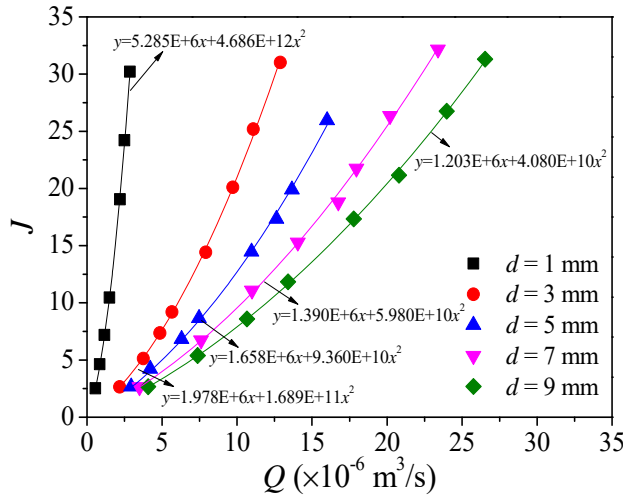
(g) G3 $\sigma_n = 1.5$ MPa



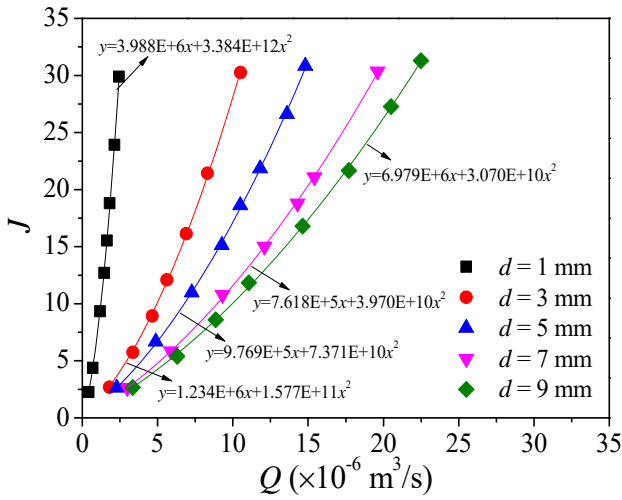
(h) G3 $\sigma_n = 2.0$ MPa



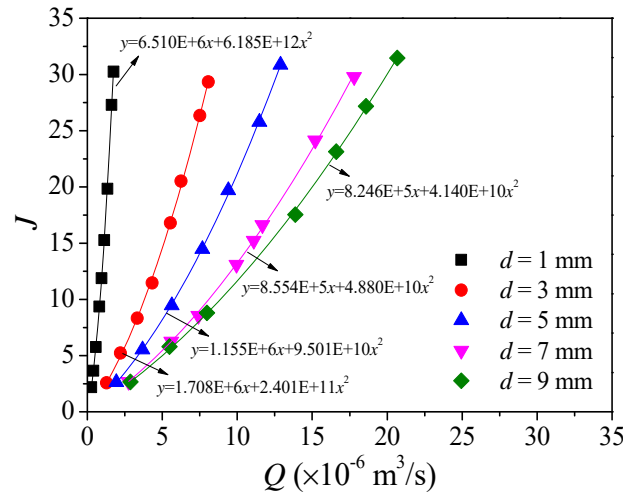
(i) G4 $\sigma_n = 0.5$ MPa



(g) G4 $\sigma_n = 1.0$ MPa



(k) G4 $\sigma_n = 1.5$ MPa



(l) G4 $\sigma_n = 2.0$ MPa

Fig. A1. Relationships between hydraulic gradient (J) and volumetric flow rate (Q) for G2-G4 under different normal stresses.